



# Atmospheric Correction of MODIS Visible and SWIR Bands: Applications to Cloud Optical Property Retrievals

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# TABLE OF CONTENTS:

- Introduction.
- Our Atmospheric Correction Method.
- A Case Study: off the coast of Peru.
- Ancillary Data Comparisons: NCEP-DAO



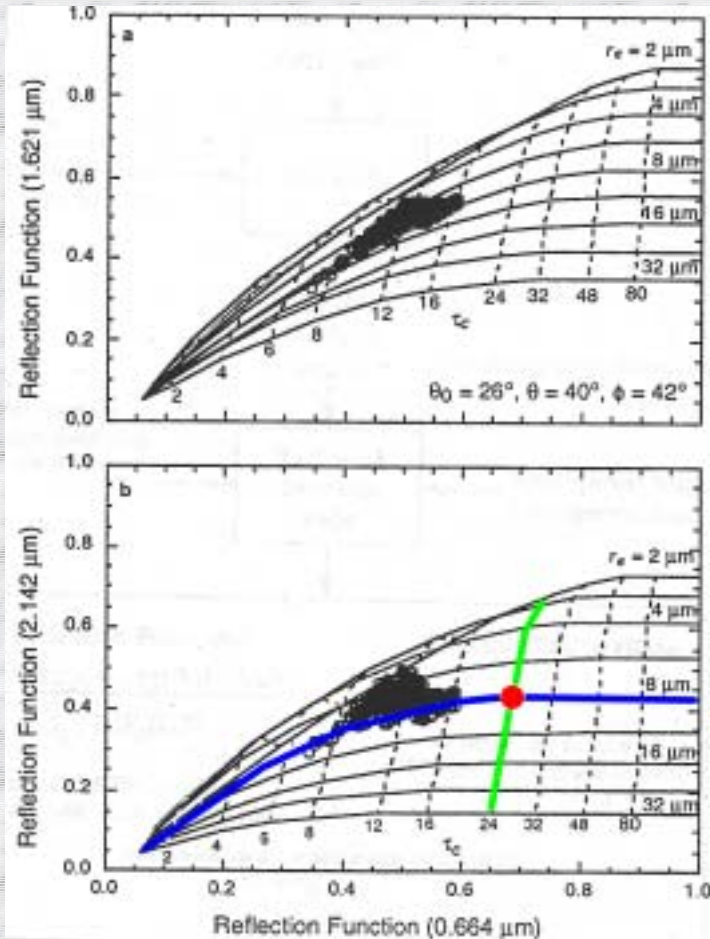
# INTRODUCTION:

## MOD06OD Algorithm in A Nut Shell

- The determination of **optical thickness ( $\tau_c$ )** and **effective radius ( $r_e$ )** is an inverse problem.
- If  $\tau_c$  and  $r_e$  are known, we can forward calculate the cloud reflectance for a given underlying surface and solar-viewing geometry, using the Mie and asymptotic theory.
- Reflectance of a non-absorbing visible band ( $0.65 \mu\text{m}$ ) is primarily function of  $\tau_c$ , while Reflectance of water/ice absorbing band ( $2.1 \mu\text{m}$ ) is primarily a function of  $r_e$ .

# INTRODUCTION:

## MOD06OD Algorithm in A Nut Shell (continued)



Values of  $\tau_c$  and  $r_e$  are obtained by comparing the  $R_{obs}$  with  $R_{calculated}$  and searching for the combination of that gives the best fit.

$$R_{obs} = \frac{\pi * I_{obs}}{\mu_0 * F_0}$$

$I_{obs}$ : observed radiance by MODIS

$F_0$ : Solar Irradiance at TOA

$\mu_0$ : cosine of solar zenith

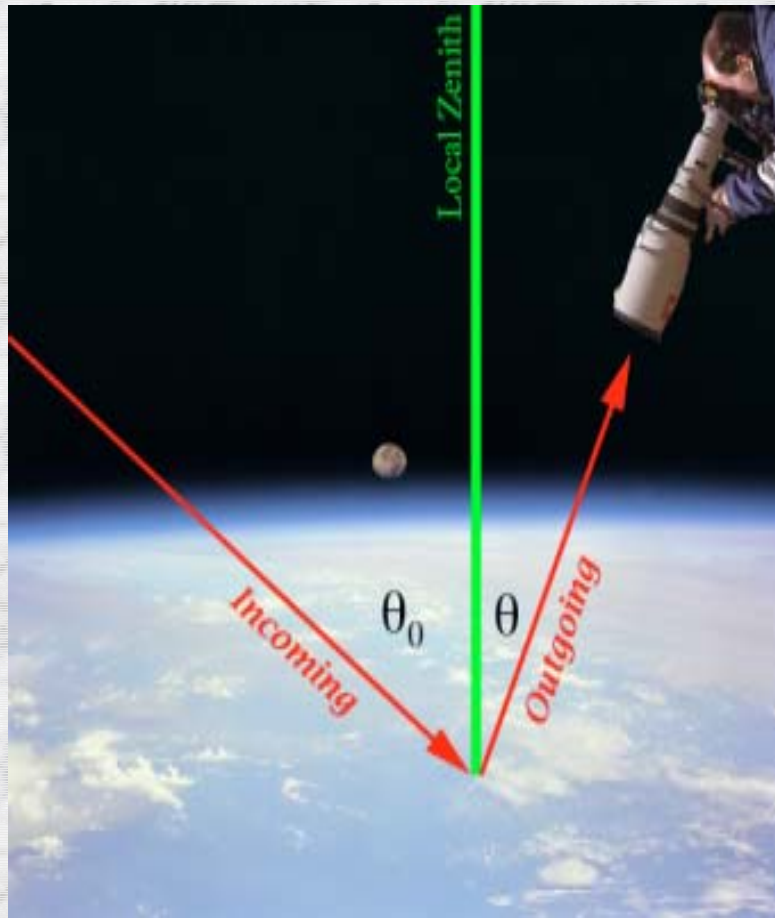
Land: 0.65 and 2.1 μm pair

Ocean: 0.86 and 2.1 μm pair



# INTRODUCTION:

## Define Problem



Monochromatic reflectance of a cloud:

$$\begin{aligned} &= \frac{\pi * I_{\lambda}^{cloudTop}}{\mu_0 * F_{\lambda}^{cloudTop}} \\ &= \frac{\pi * \frac{I_{\lambda}^{obs}}{T_{\lambda}(\mu)}}{\mu_0 * F_{0\lambda} * T_{\lambda}(\mu_0)} \\ &= \frac{\pi * I_{\lambda}^{obs}}{\mu_0 * F_{0\lambda}} * \frac{1}{T_{\lambda}(\mu) * T_{\lambda}(\mu_0)} \end{aligned}$$

# INTRODUCTION:

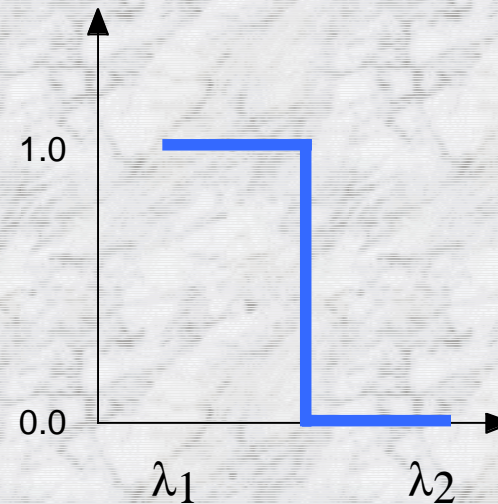
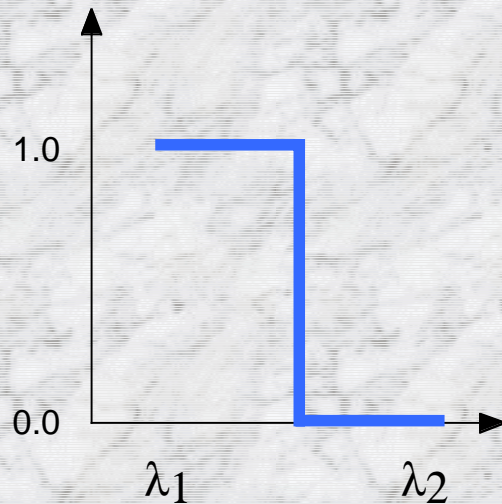
## Define Problem

Two-way transmittance for a MODIS Band:

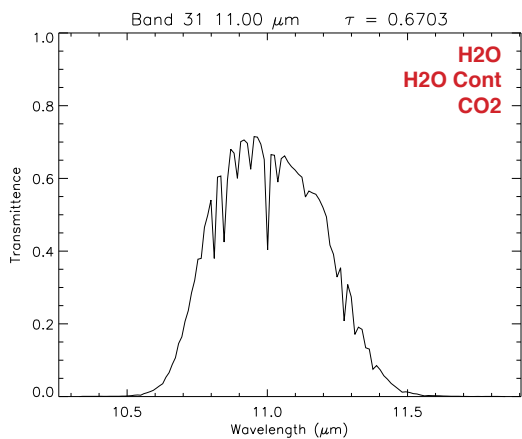
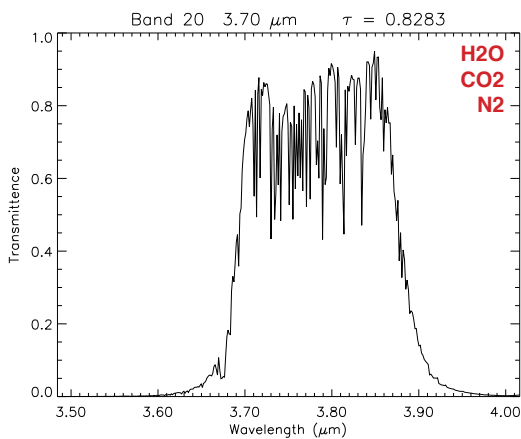
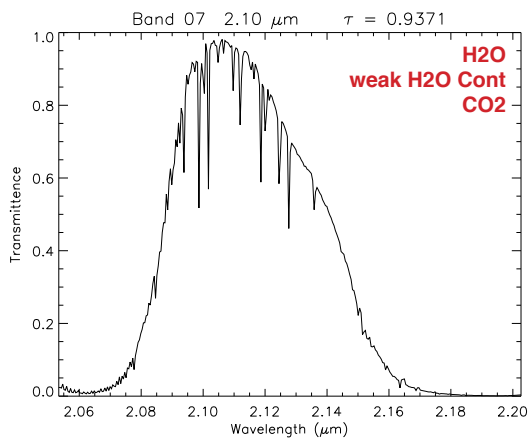
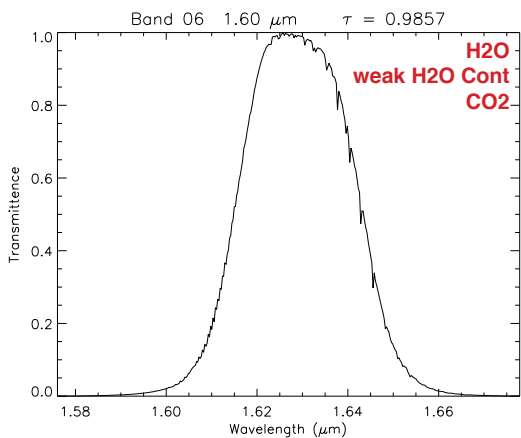
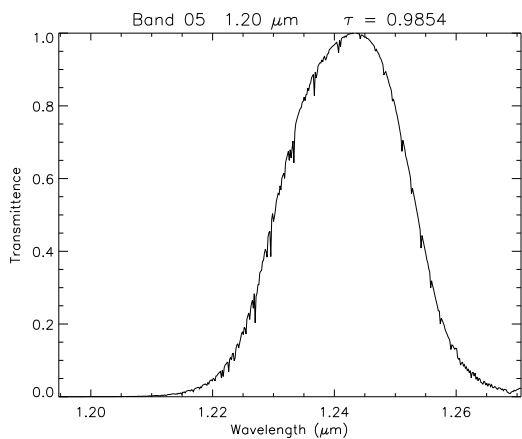
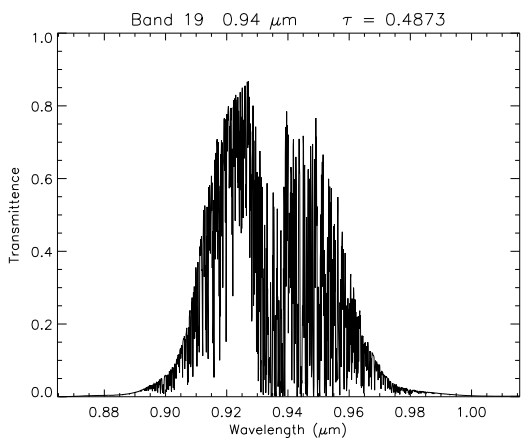
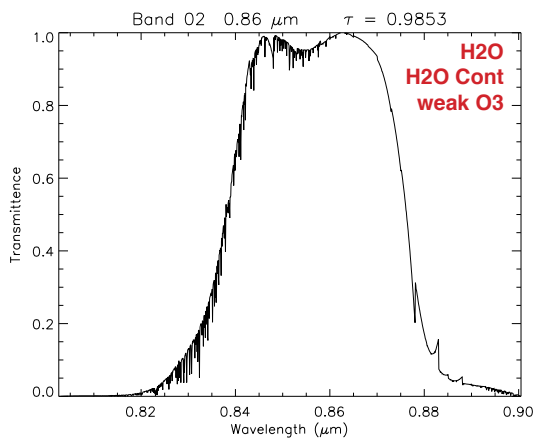
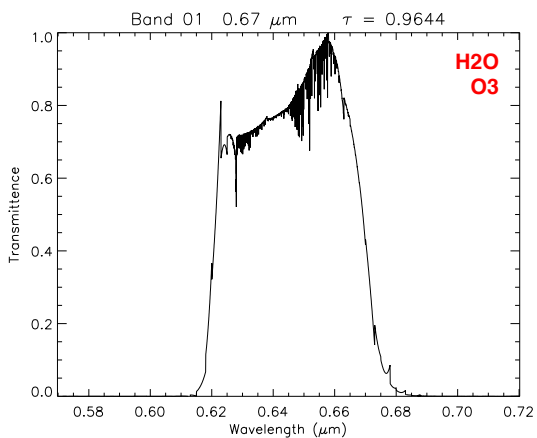
$$[T(\mu_0) * T(\mu)]_{band} = \frac{1}{2} \neq [T(\mu_0)]_{band} * [T(\mu)]_{band}$$

$$T(\mu_0) = 1/2$$

$$T(\mu) = 1/2$$



# Mid-Latitude Summer Model Atmosphere (no Rayleigh)





# CORRECTION METHOD:

## Parameterization

- Column integrated precipitable water from cloud top to TOA. Other common gases included, although not parameterized.
- *Moisture weighted mean column air temperature from cloud top to TOA.*
- *Surface temperature.*



# CORRECTION METHOD:

## Parameterization (continued)

Transmittance( $P_c$ ,  $PW$ ,  $\mu_{\text{effective}}$ ,  $\lambda$ ):

■  $P_c$  = cloud top height

■  $PW$  = precipitable water from cloud top to TOA

■  $\mu_{\text{effective}}$  is defined as: 
$$\frac{1}{\mu_{\text{effective}}} = \frac{1}{\mu_0} + \frac{1}{\mu}$$

■  $\lambda$ : wavelength

# CORRECTION METHOD: MODTRAN 4

- For each model atmosphere, compute **one-way** transmittances for various combinations of  $P_c$ ,  $P_W$ ,  $\mu_{\text{effective}}$ . The result is a 3D table:  $\mathbf{T}(P_c, P_W, \mu_{\text{effective}})$ .
- Repeat the same procedure for all available model atmospheres, plus several real-world atmospheric profiles to increase sampling space. Now you have a 4D table:  $\mathbf{T}(P_c, P_W, \mu_{\text{effective}}, N_{\text{Profiles}})$
- Collapse 4th dimension by taking averages, also store standard deviations:  $\mathbf{T}_{\text{ave}}(P_c, P_W, \mu_{\text{effective}})$ .



# CORRECTION METHOD: MODTRAN 4 (continued)

- The linkage between one-way and two-way transmittance is in  $\frac{1}{\mu_{effective}}$ , the path (i.e. absorbing amount scaling factor)

$$\frac{1}{\mu_{effective}} = \frac{1}{\cos \theta_{effective}} = \frac{1}{\cos \theta_0} + \frac{1}{\cos \theta}$$

- If it is a two-way path,  $\mu_{effective} \leq 0.5$  (or  $\theta_{effective} \geq 60^\circ$ ).
- A MODIS granual:  $\theta_0 = 33.32^\circ$ ,  $\theta = 46.16^\circ$  (edge pixel), then  $\theta_{effective} = 67.75^\circ$ .

# CORRECTION METHOD:

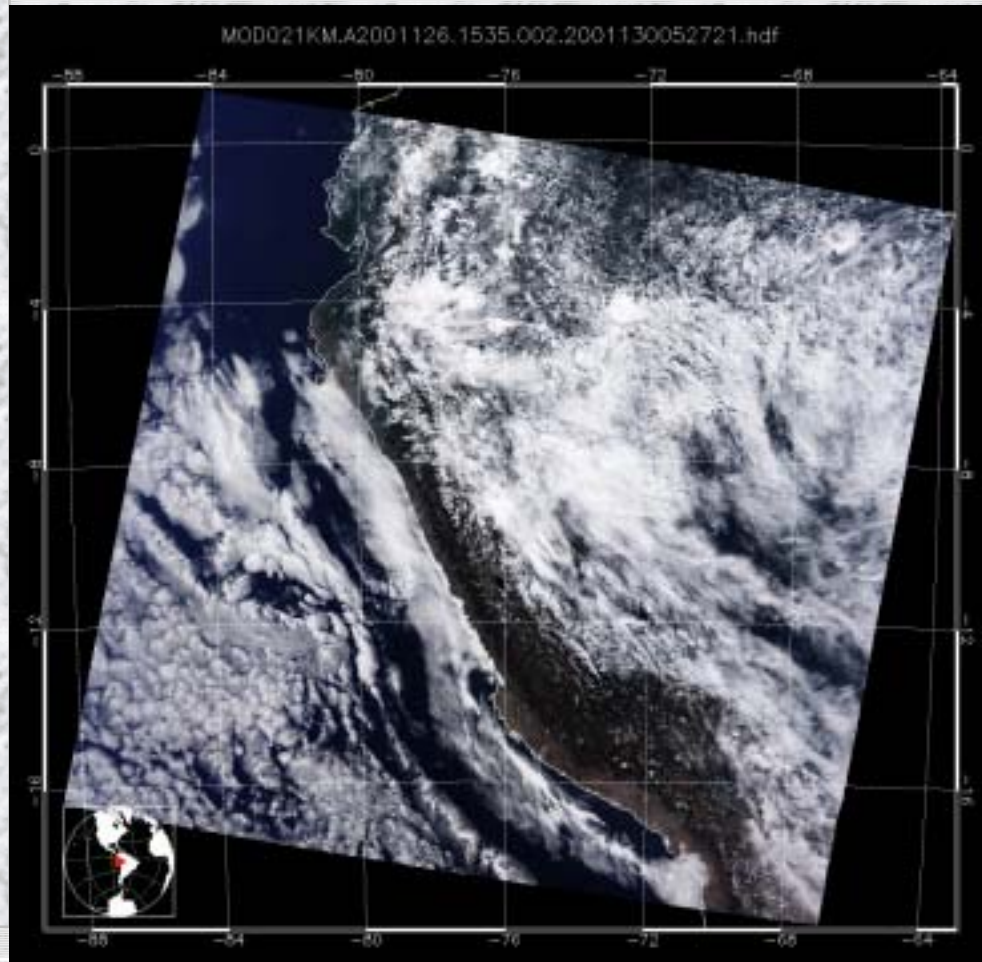
## Sources of Uncertainties

- Refraction is not properly accounted for.
- Not a 3-D geometry.
- Variations in the structures of moisture profiles.



# CASE STUDY:

## RGB Composite



### Date:

May 6<sup>th</sup>, 2001  
126<sup>th</sup> day

### Time:

15:35 UTC  
10:35 Local

### Ecosystems:

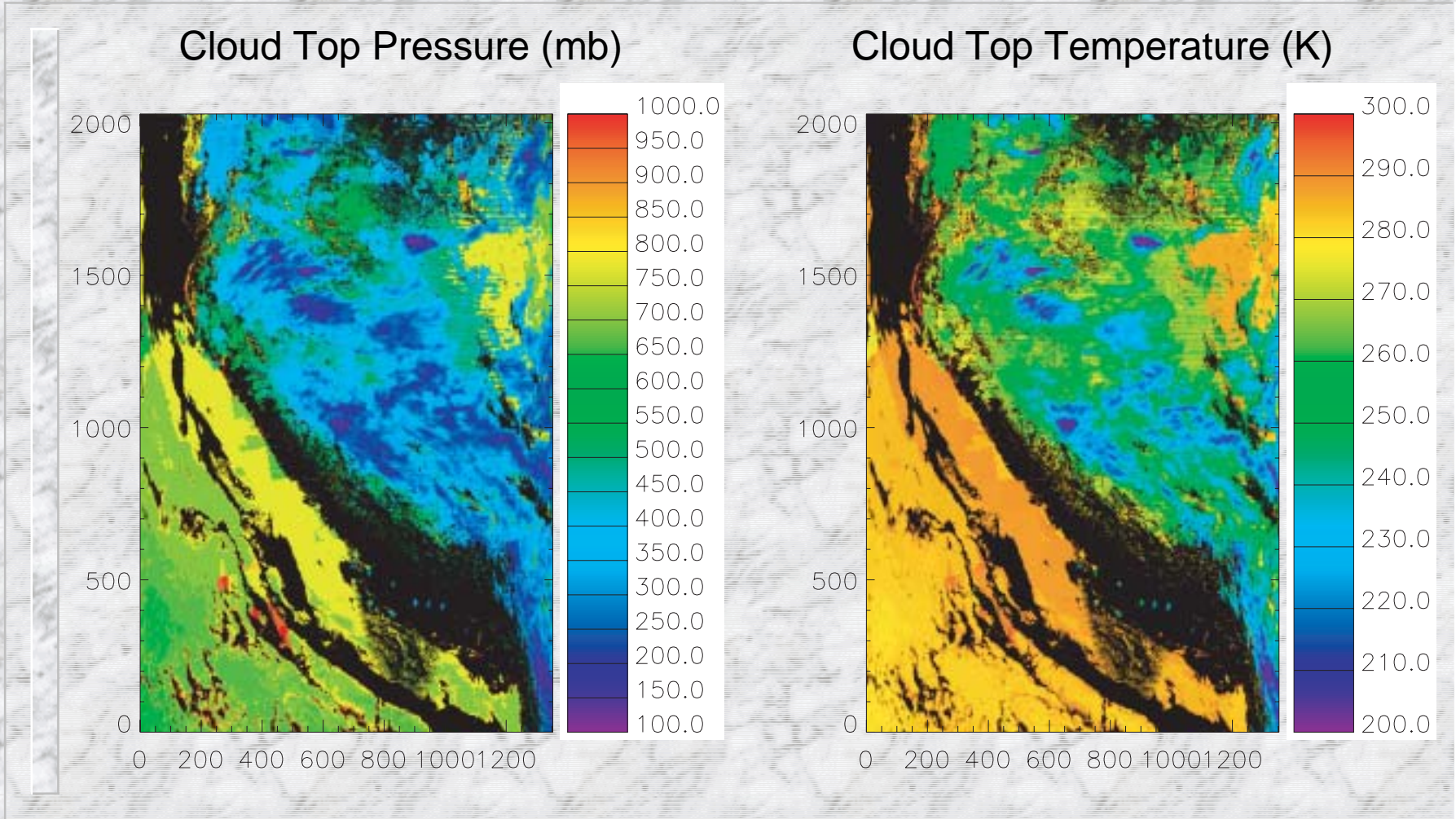
Ocean, narrow strip of  
grass land, evergreen  
forest.

# SOUTH AMERICA





# Case Study: Cloud Top Pressure and Temperature



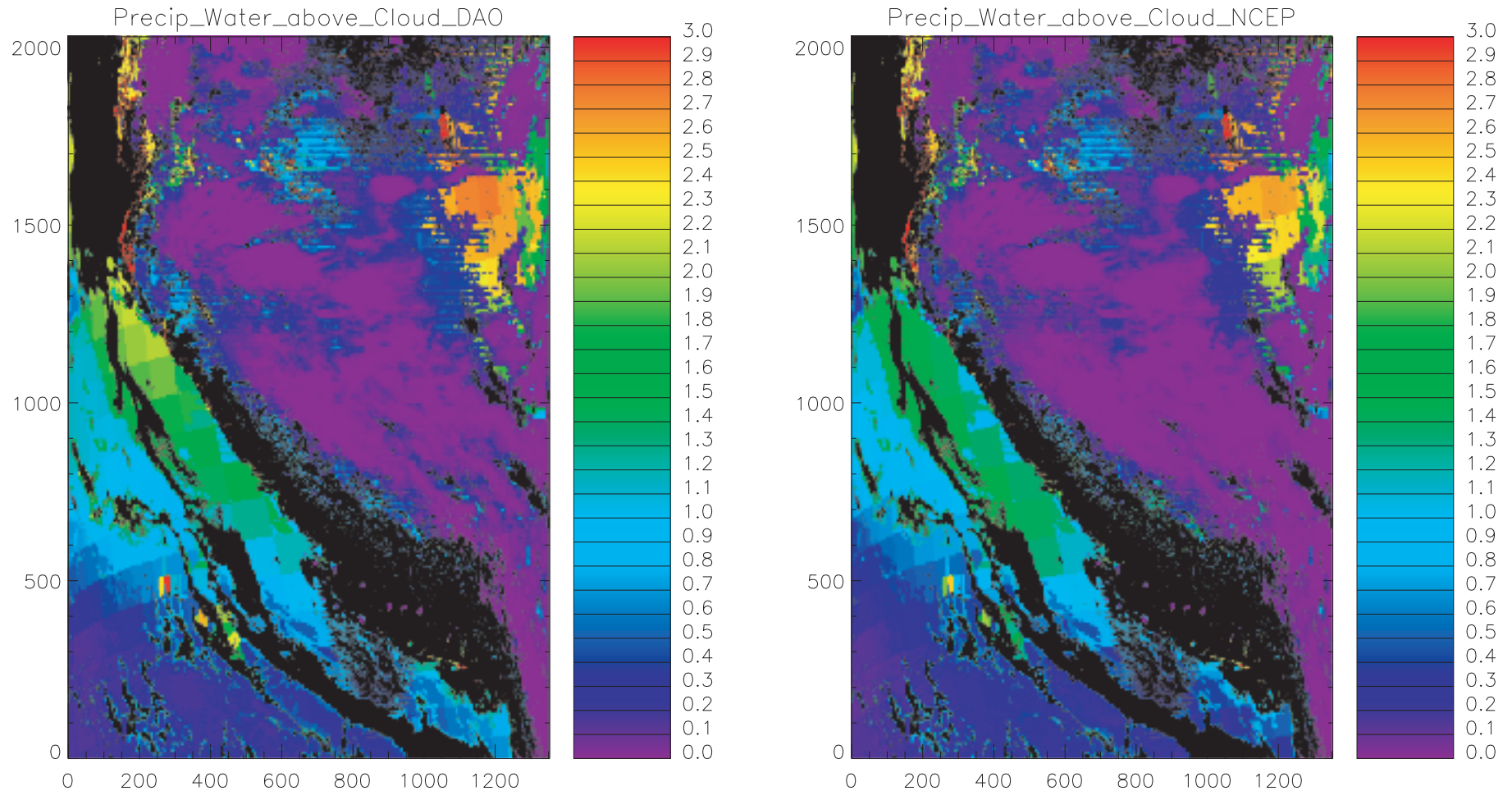
January 8, 2002

Case Study

15

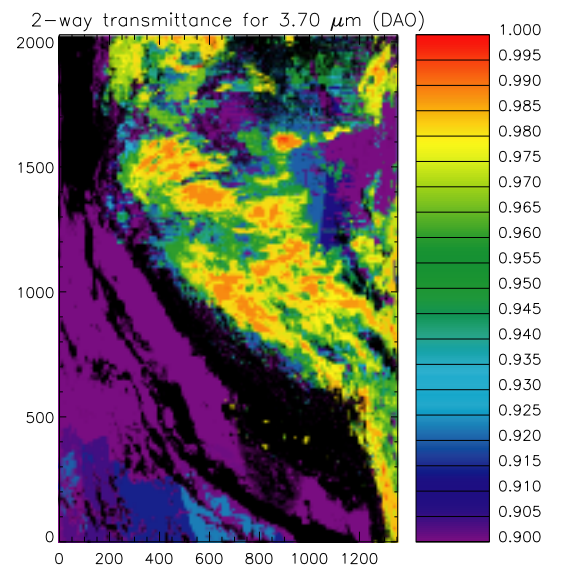
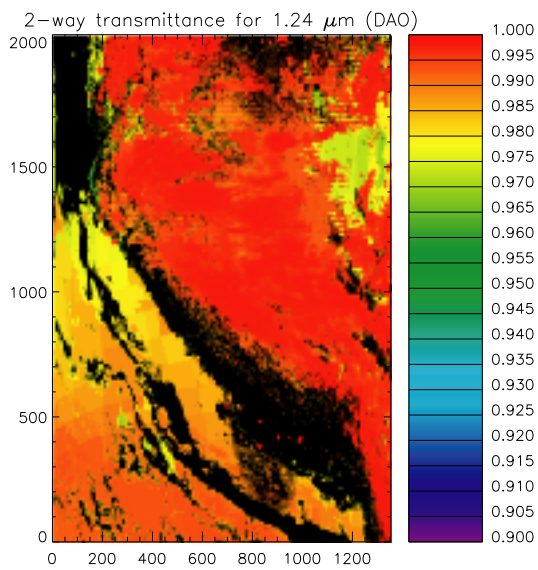
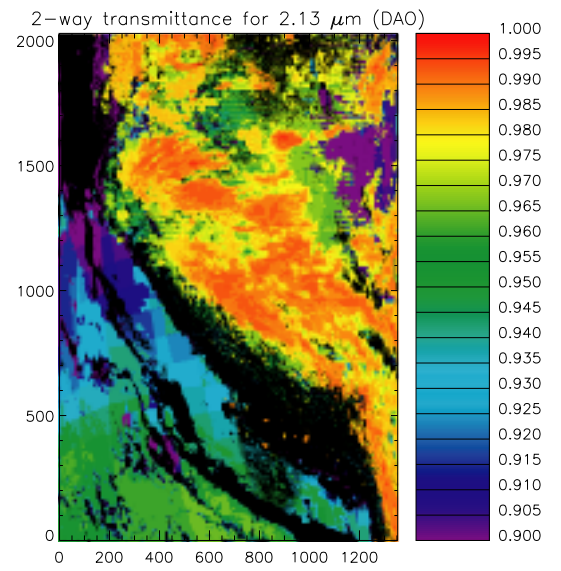
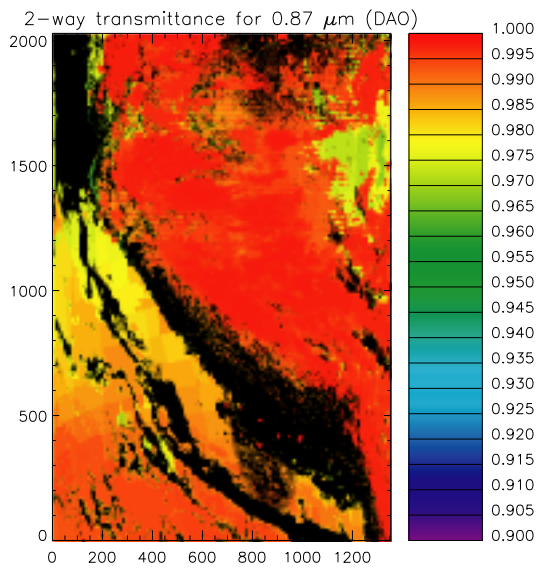
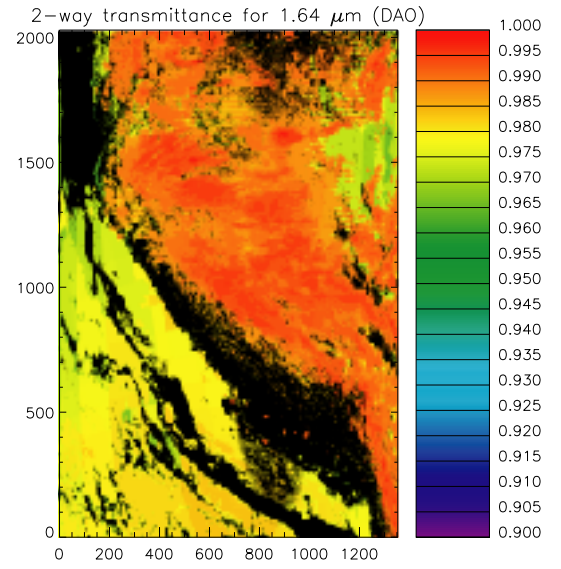
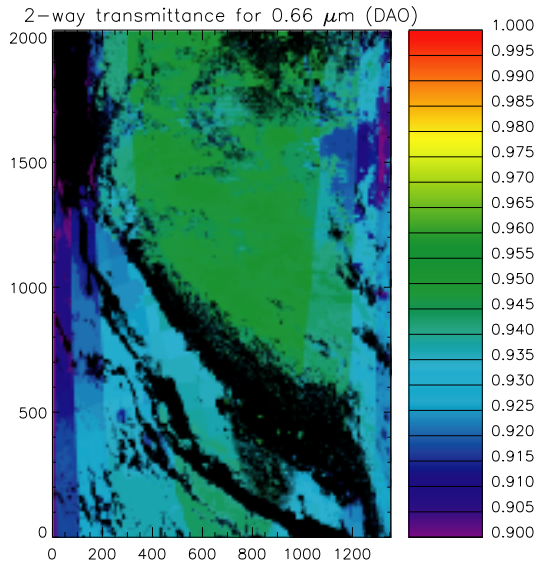
# Case Study

## Integrated Precipitable Water Above Cloud



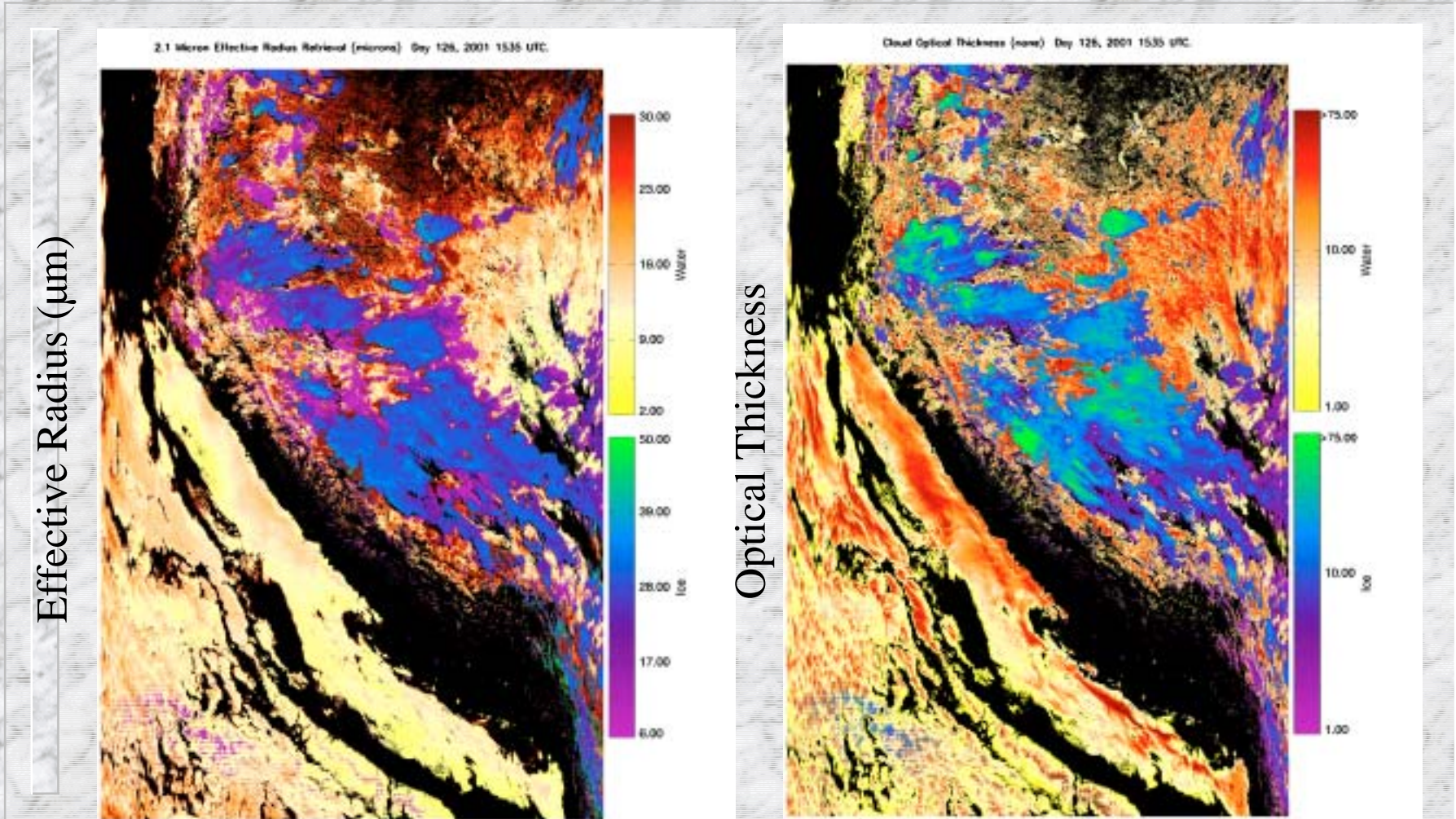


# Two-Way Transmittance (DAO)





# Case Study: Retrieval Results (atmospheric corrected)



Effective Radius ( $\mu\text{m}$ )

Optical Thickness

January 8, 2002

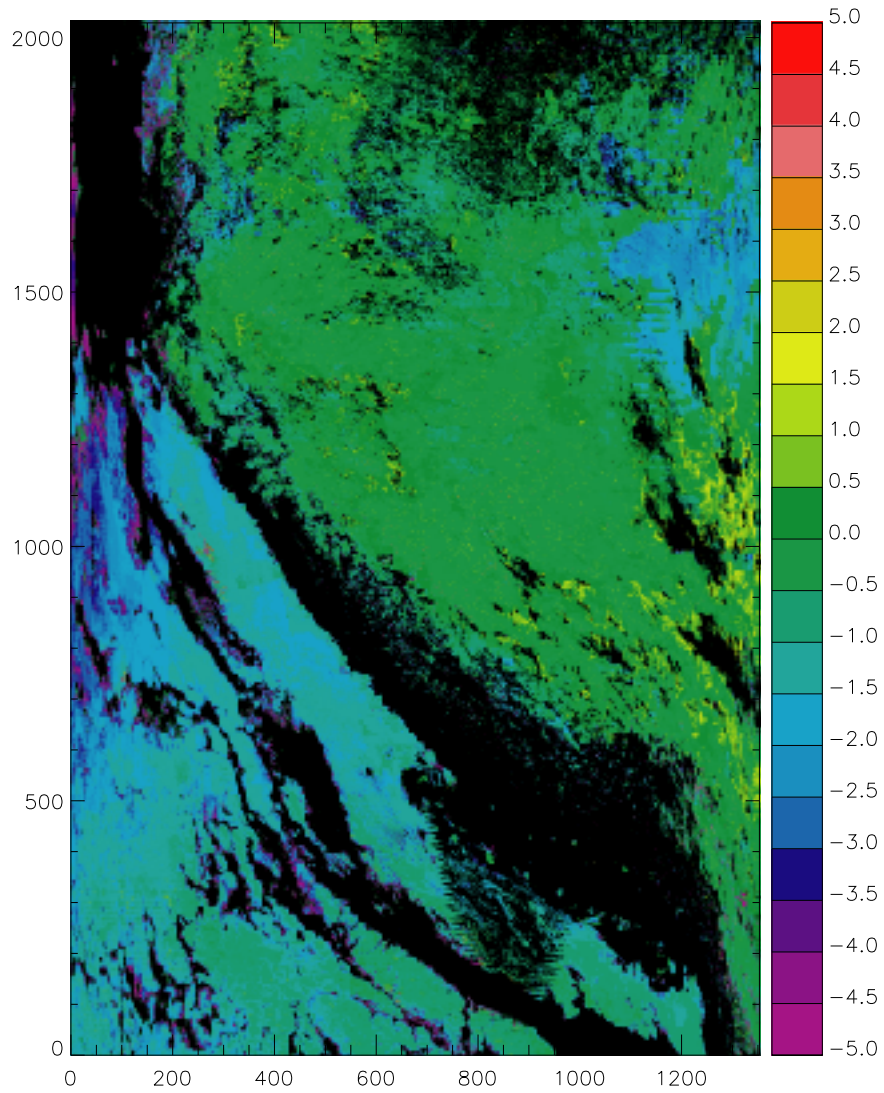
Case Study

18

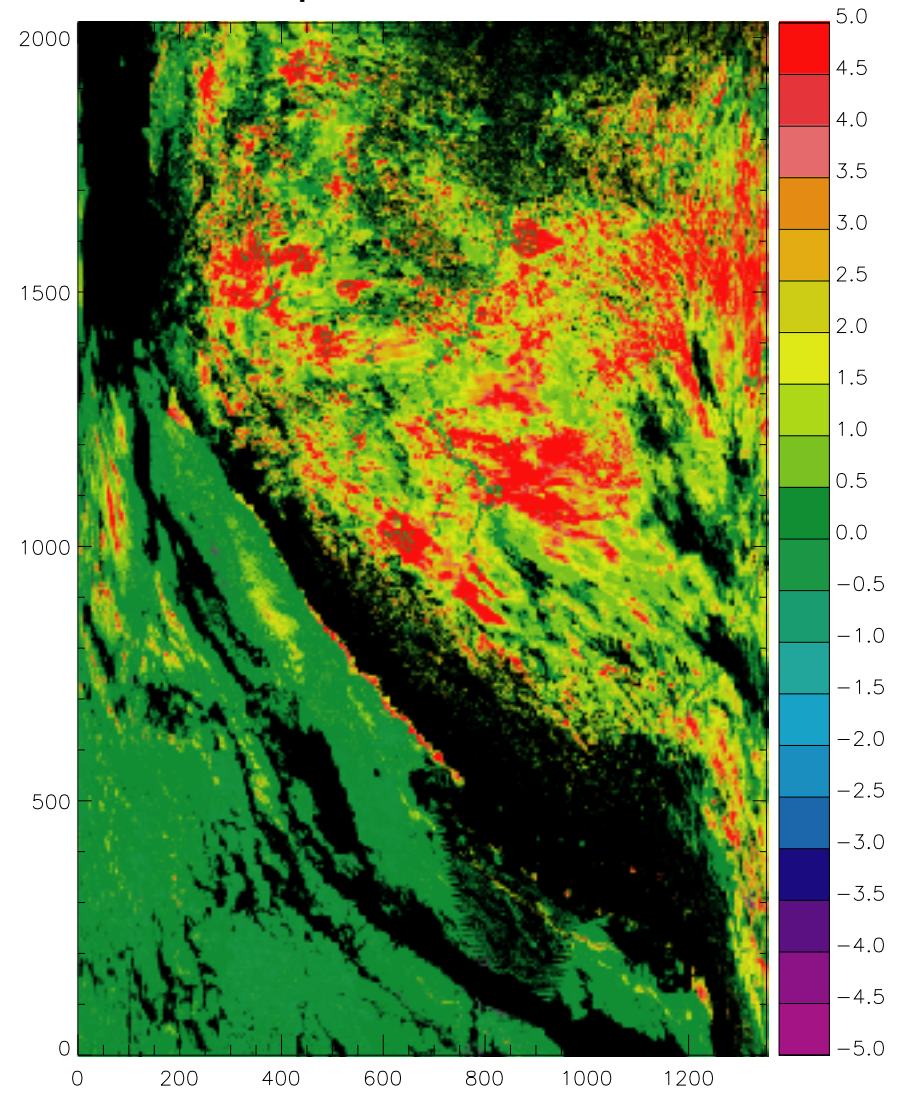


# Differences (atmospheric corrected - uncorrected)

## Effective Radius ( $\mu\text{m}$ )



## Optical Thickness



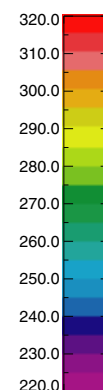
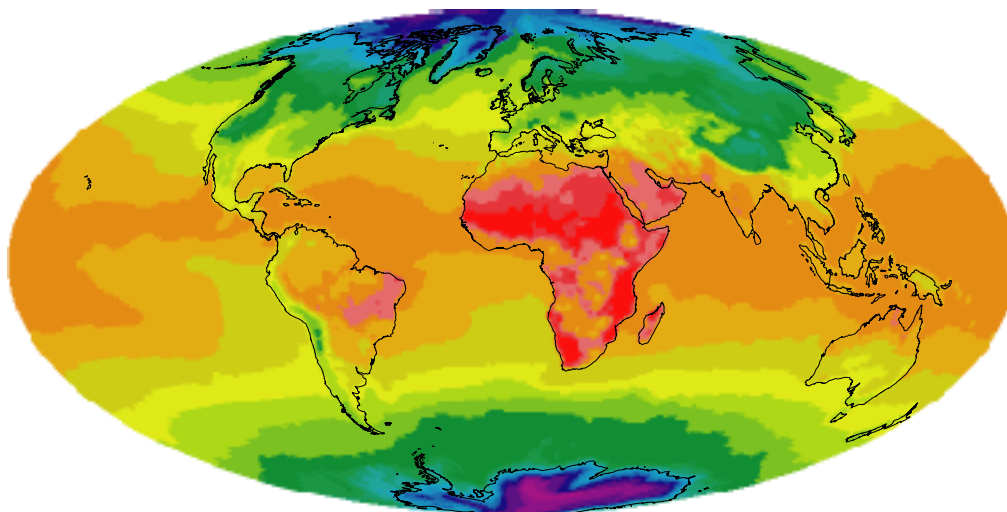
# Ancillary Data Comparison

DAO versus NCEP

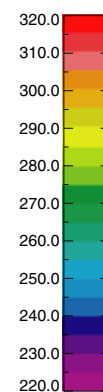
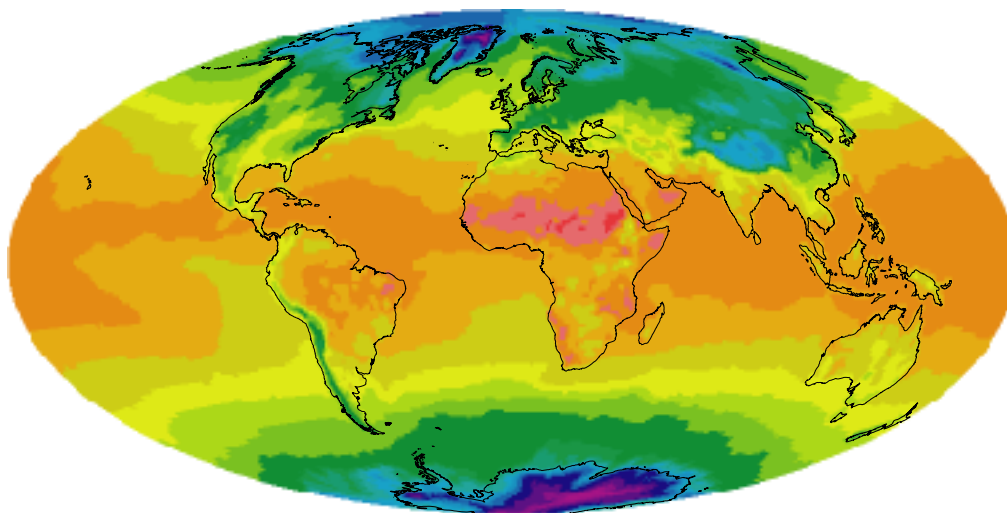


# DAO and NCEP Surface Temperature Comparisons November 14, 2001 12:00 UTC

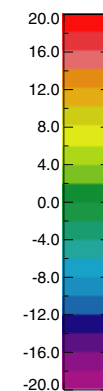
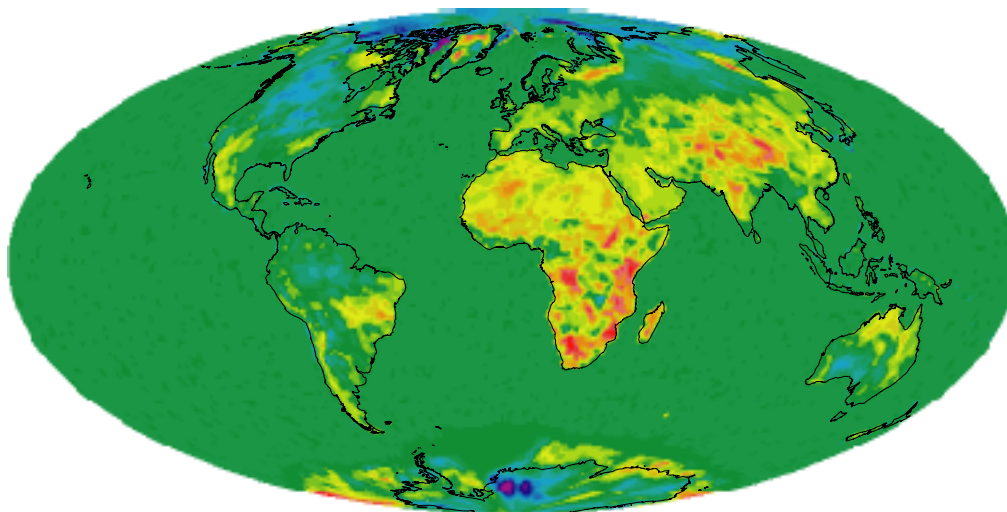
DAO



NCEP

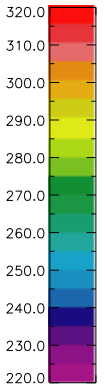
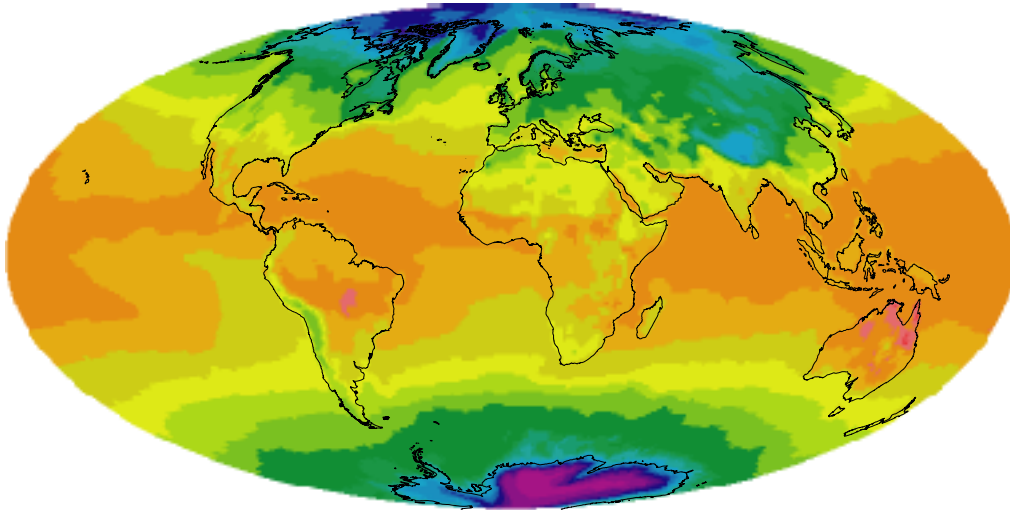


DAO - NCEP

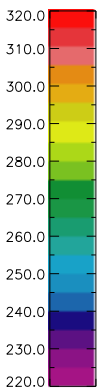
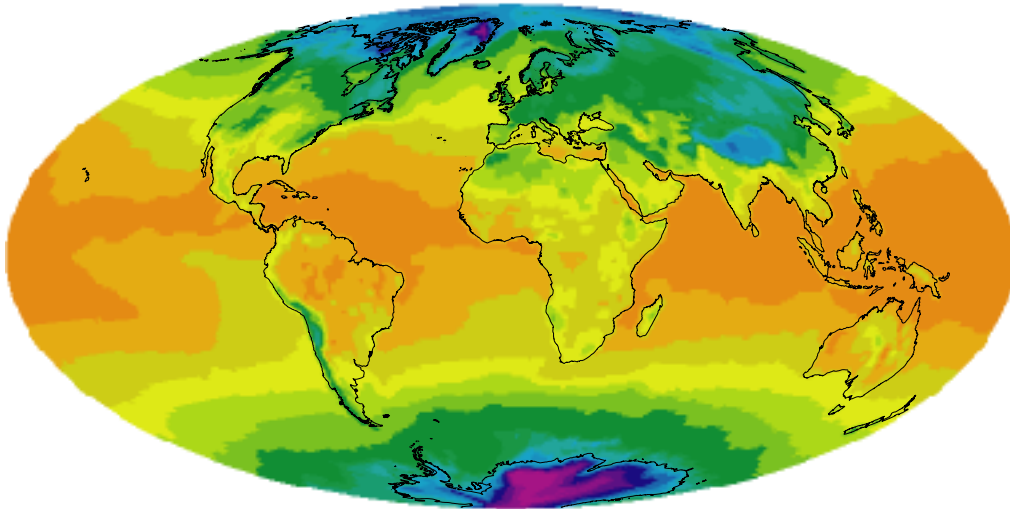


# DAO and NCEP Surface Temperature Comparisons November 14, 2001 00:00 UTC

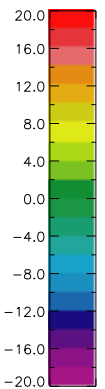
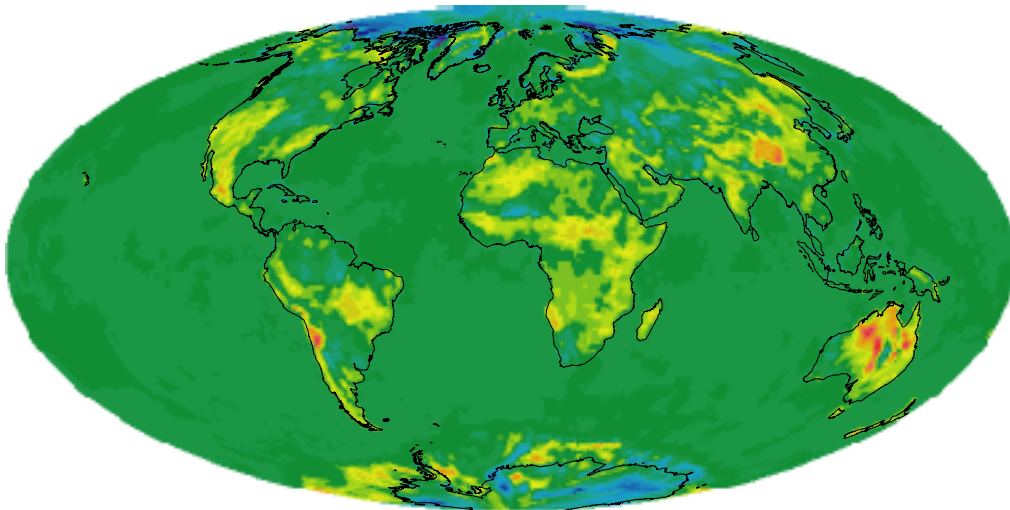
DAO



NCEP



DAO - NCEP

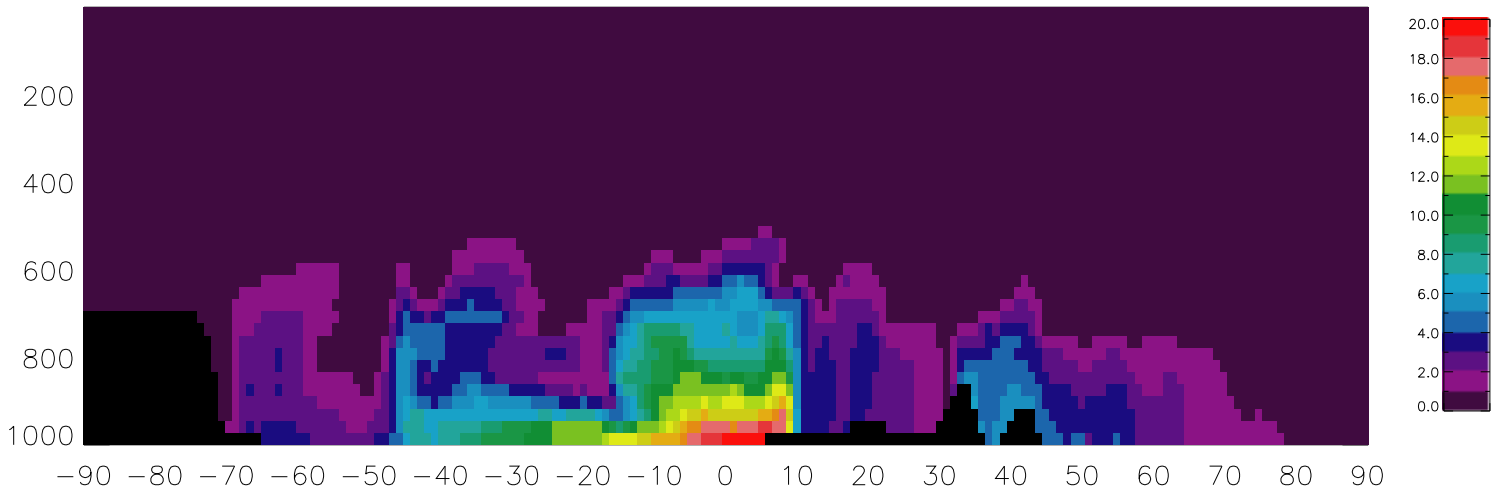




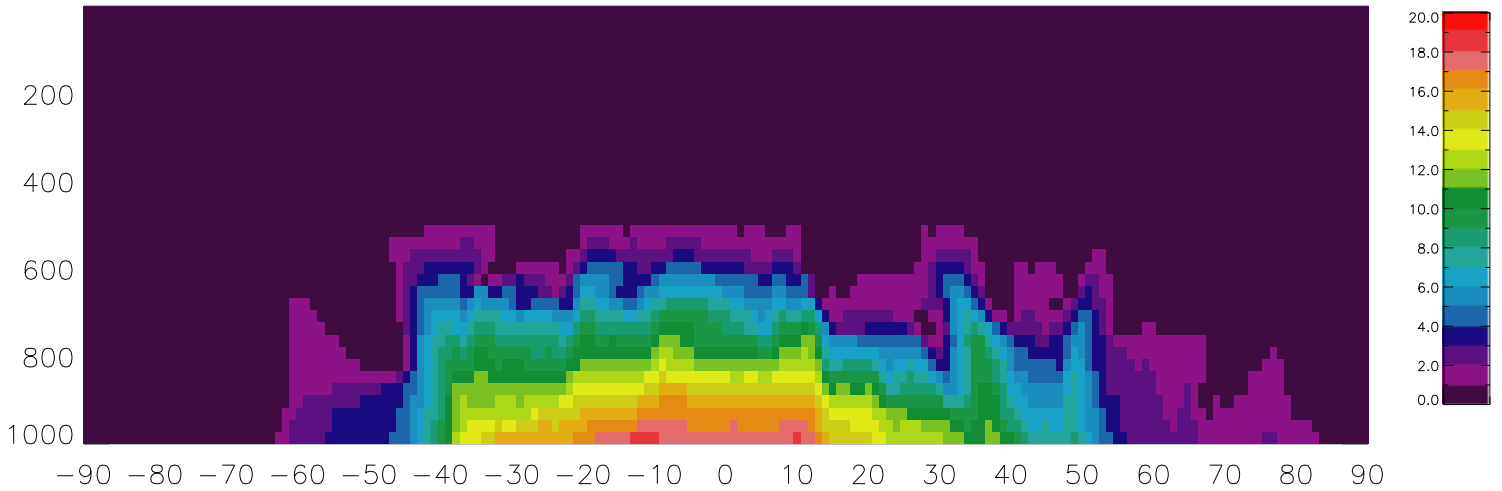
# Latitude-Height Cross Section of Specific Humidity (g/kg)

Longitude = 0 degree, November 14, 2001, 00:00UTC

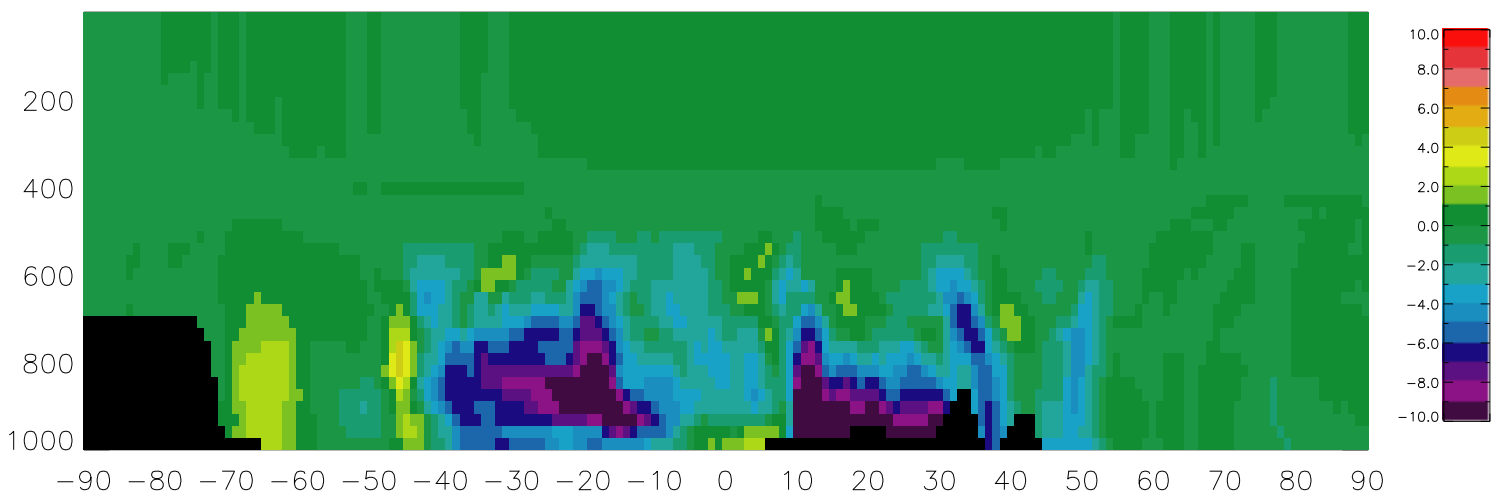
DAO



NCEP



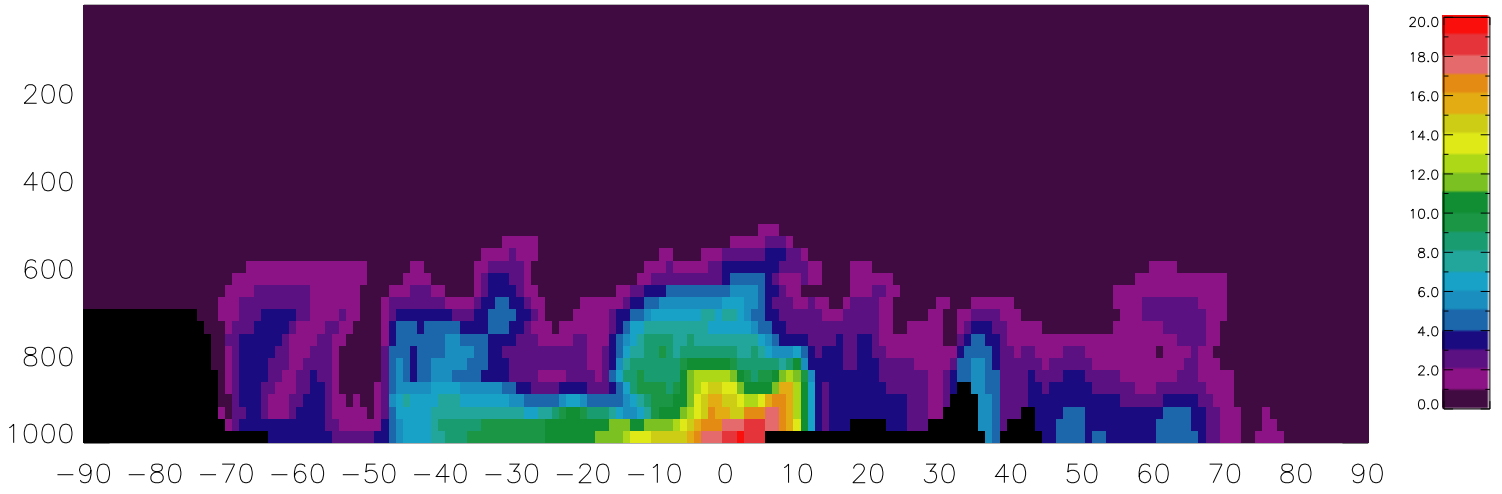
DAO - NCEP



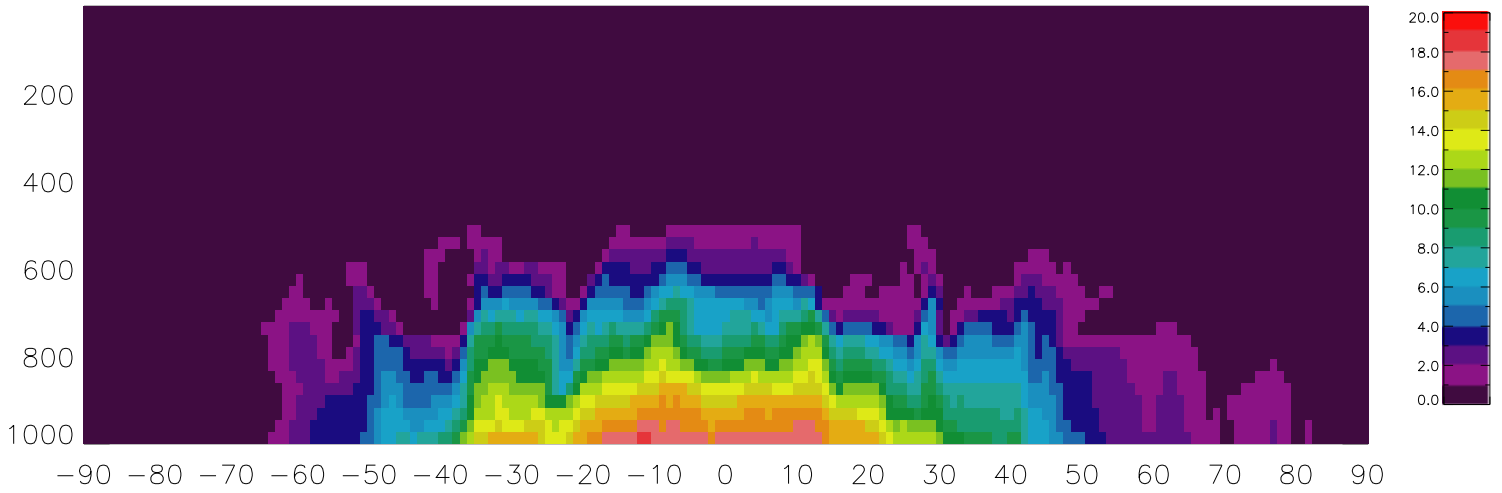
# Latitude-Height Cross Section of Specific Humidity (g/kg)

Longitude = 0 degree, November 14, 2001, 12:00UTC

DAO



NCEP



DAO - NCEP

