

An Algorithm for Coastal Water and the Status of its Implementation into the MODIS Processing Stream

by

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CASE 1 WATERS			
1	LIVING ALGAL CELLS <i>variable concentration</i>	RESUSPENDED SEDIMENTS	4 <i>from bottom along the coast- line and in shallow areas</i>
2	ASSOCIATE DEBRIS <i>originating from grazing by zooplankton and natural decay</i>	TERRIGENOUS PARTICLES	5 <i>river and glacial runoff</i>
3	DISSOLVED ORGANIC MATTER <i>liberated by algae and their debris (yellow substance)</i>	DISSOLVED ORGANIC MATTER	6 <i>land drainage (terrigenous yellow substance)</i>
		ANTHROPOGENIC INFLUX	7 <i>particulate and dissolved materials</i>
CASE 2 WATERS			

(Gordon and Morel, 1983)

Atmospheric Correction

$$\rho_t(\lambda) = \rho_r(\lambda) + \rho_A(\lambda) + t_v(\lambda)t_s(\lambda)\rho_w(\lambda)$$

Case 1 waters :

- $\rho_w(765) \approx \rho_w(865) \approx 0$, \Rightarrow NIR can be used to assess the aerosol influence.

Case 2 waters :

- $\rho_w(\text{NIR}) \neq 0 \Rightarrow$ no bands "tailor made" for assessing the aerosol.
- Case 2 waters contain large quantities of dissolved organic material that influence ρ_t in a manner similar to strongly absorbing aerosols.
- Strongly absorbing aerosols are often found near the coast.

Approach for Case 2 waters: model $\rho_A(\lambda)$ and $\rho_w(\lambda)$, and then use spectral optimization to find the best values of the model parameters.

The Aerosol Model

Uses a Junge Power -Law Size Distribution :

$$\frac{dN}{dD} \equiv \frac{K}{D^{\nu+1}} \quad D < D_0,$$
$$\frac{dN}{dD} = 0$$

$$D_0 \leq D \leq D_1,$$

$$D_1 \leq D \leq D_2,$$

$$D > D_2,$$

$D_0 = 0.06 \mu\text{m}$, $D_1 = 0.20 \mu\text{m}$, and $D_2 = 20 \mu\text{m}$.

Mie theory is used to compute aerosol properties

- $m = m_r - im_i$, where m_r is either 1.50 or 1.333, and $m_i = 0, 0.001, 0.003, 0.010, 0.030, \text{ and } 0.040$.

$$\rho_p(G, \lambda, m_r, m_i, v) = a(G, \lambda, m_r, m_i, v) \tau(\lambda) + b(G, \lambda, m_r, m_i, v) \tau^2(\lambda) + c(G, \lambda, m_r, m_i, v) \tau^3(\lambda) + d(G, \lambda, m_r, m_i, v) \tau^4(\lambda)$$
- v ranges from 2.0 to 4.5 in steps of 0.5.
- 72 separate aerosol models (2 values of $m_r \times 6$ values of $m_i \times 6$ values of v).
- Interpolate to essentially give a continuum of models.

The Water Model (Garver and Siegel, 1997)

$$\rho_w = \rho_w(b_b/(a+b_b))$$

$$a = a_w + a_{ph} + a_{cdm}$$

$$b_b = b_{bw} + b_{bp}$$

$$a_{ph}(\lambda) = a_{ph0}(\lambda) C$$

$$a_{cdm}(\lambda) = a_{cdm}(443) \exp[-S(\lambda-443)]$$

$$b_{bp}(\lambda) = b_b(443) [443/\lambda]^n$$

$$\rho_w = \rho_w(\lambda, C, a_{cdm}(443), b_{bp}(443))$$

Note, the parameters $a_{ph0}(\lambda)$, S , and n are provided by fitting the model to experimental data. For Case 1 waters, $S = 0.0206 \text{ nm}^{-1}$ and $n = 1.03$ (Maritorena, *et al.*, 2002).

The Optimization

$$\begin{aligned} \hat{\rho}_{Aw}(G, \lambda, m_r, m_i, \nu, \tau_a, C, a_{cdm}(443), b_{bp}(443)) &\equiv \hat{\rho}_A(G, \lambda, m_r, m_i, \nu, \tau_a) \\ &+ \hat{t}_\nu(G, \lambda, m_r, m_i, \nu, \tau_a) \hat{t}_s(G, \lambda, m_r, m_i, \nu, \tau_a) \hat{\rho}_w(\lambda, C, a_{cdm}(443), b_{bp}(443)). \\ \rho_{Aw}(G, \lambda, measured) &\equiv \rho_A(G, \lambda) + t_\nu(G, \lambda) t_s(G, \lambda) \rho_w(\lambda) \end{aligned}$$

$$\rho_t(\lambda) = \rho_r(\lambda) + \rho_A(\lambda) + t_\nu(\lambda) t_s(\lambda) \rho_w(\lambda)$$

The modeled counterpart of

Assuming $\rho_A(765)$ and $\rho_A(865) = 0$ gives estimation of the parameters v and τ_a as functions of m_r and m_i , i.e., $v(m_r, m_i)$ and $\tau_a(m_r, m_i)$.

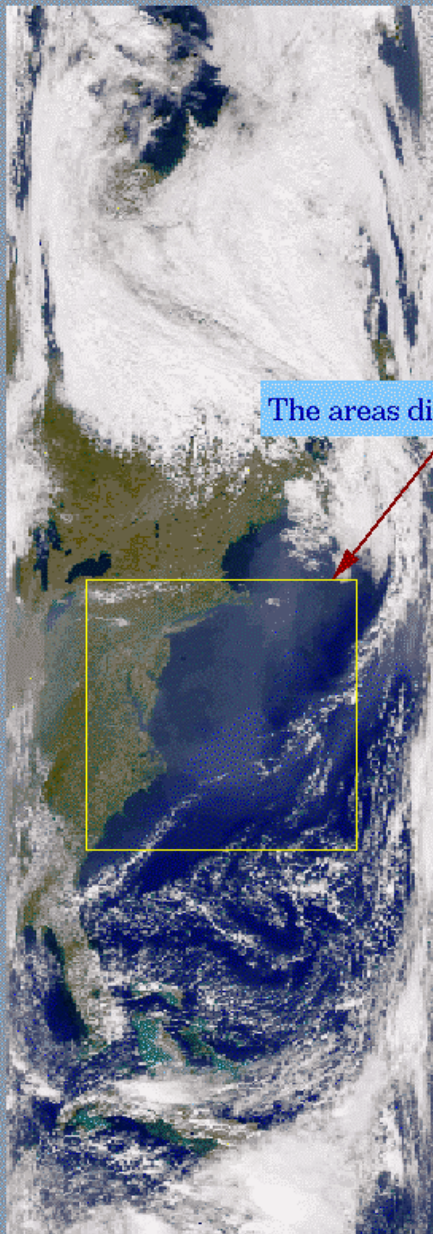
Given the constraints $v(m_r, m_i)$ and $\tau_a(m_r, m_i)$ we minimize the quantity

$$\sum \left\{ \hat{\rho}_{Aw}(G, \lambda_i, m_r, m_i, v, \tau_a, C, a_{cdm}(443), b_{bp}(443)) - \rho_{Aw}(G, \lambda_i, measured) \right\}^2$$

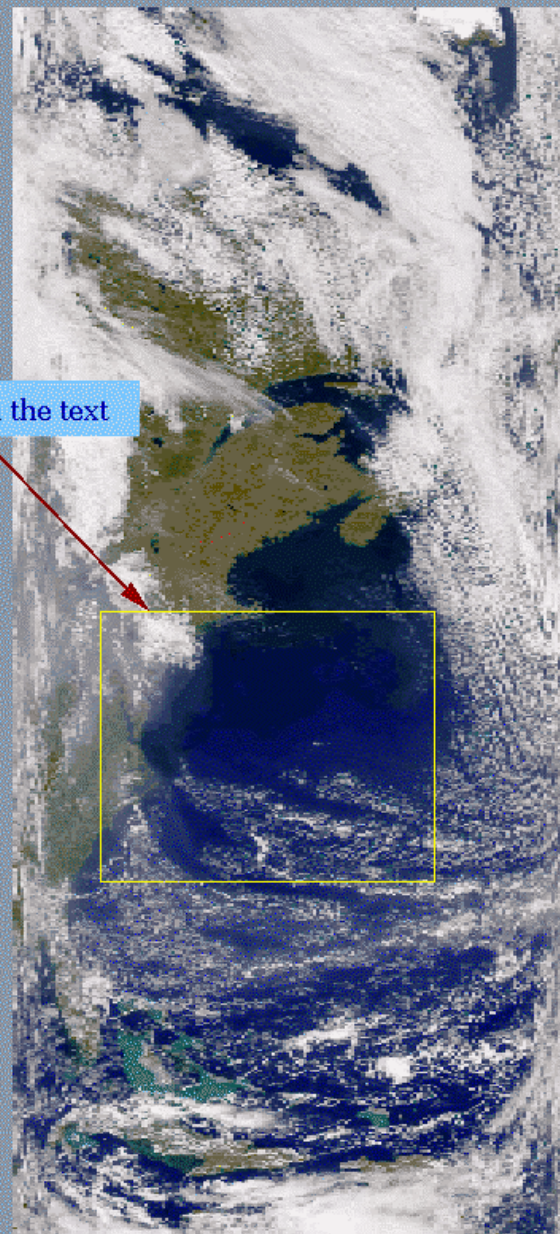
In effect, we have optimized for 7 parameters:

$C, a_{cdm}(443), b_{bp}(443), v, \tau_a, m_r,$ and m_i ;

This is generally all that is needed in Case 1 waters.



S1997279171919.L1A_HNSG_BRS
DAY - 279, October 06, 1997

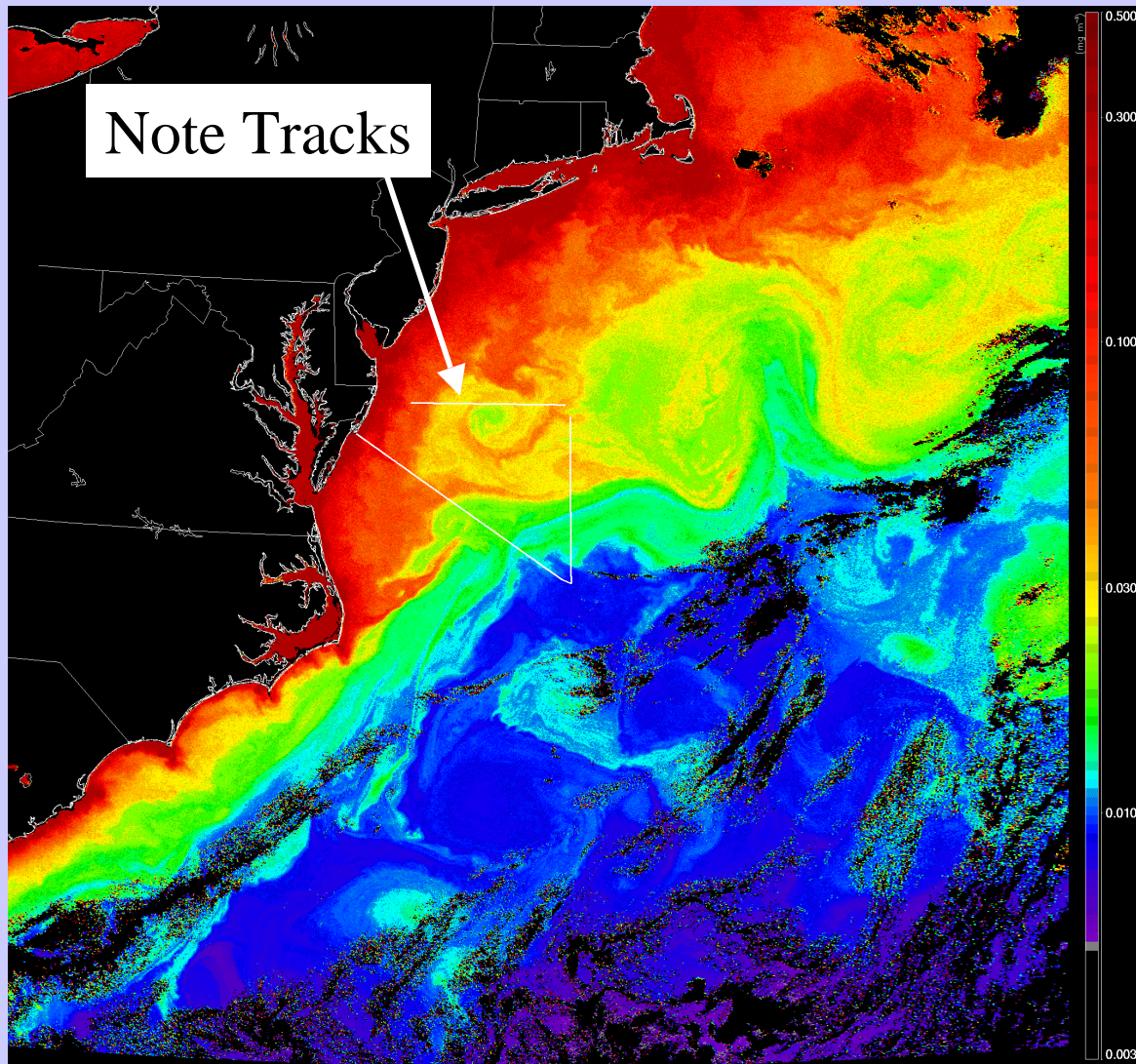


S1997281170357.L1A_HNSG_BRS
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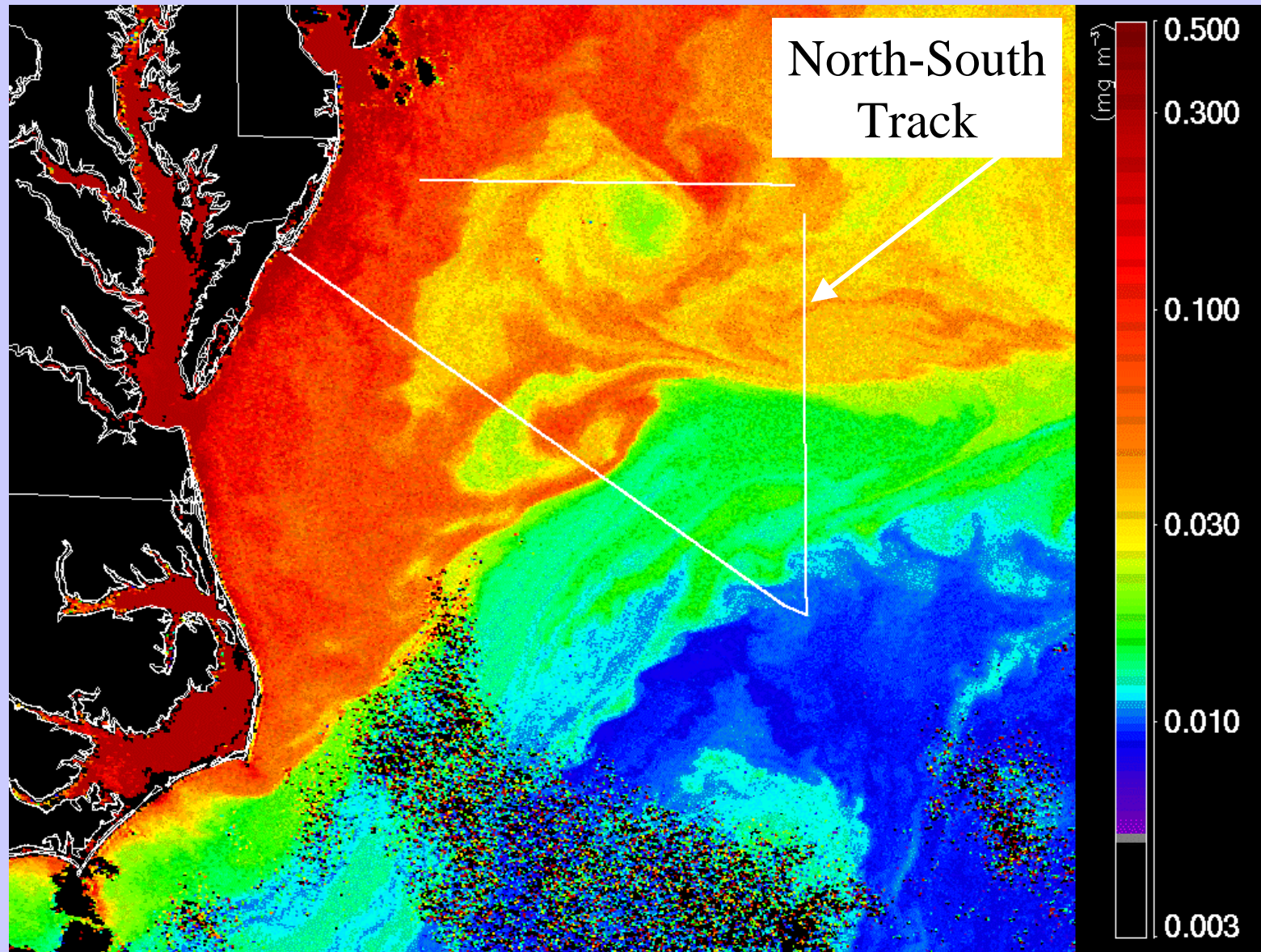
The areas discussed in the text

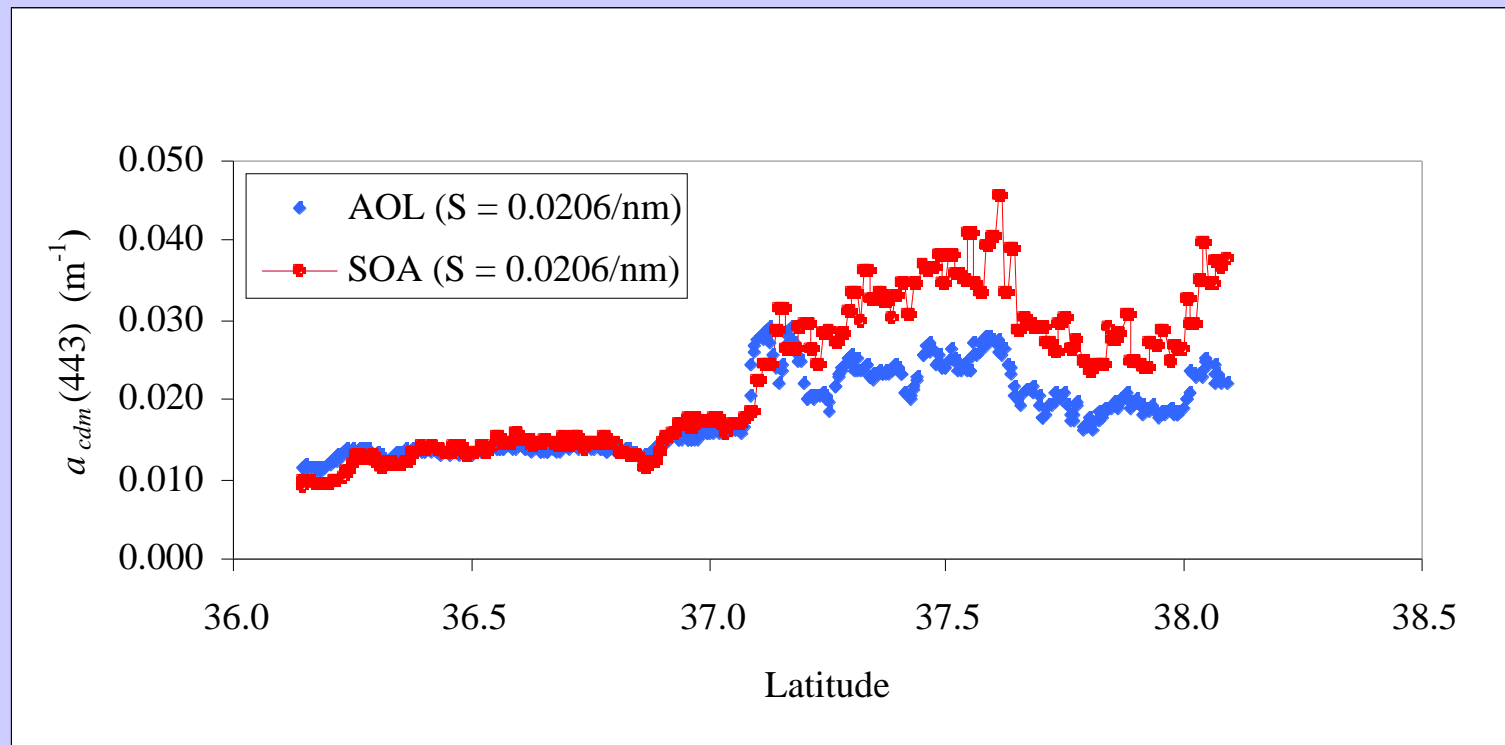
To validate this algorithm, we use the SeaWiFS image from Day 279 (left on previous slide) and compare the retrievals of a_{cdm} from the algorithm with estimates of a_{CDOM} from the AOL. The AOL measurements are made along the triangular path drawn on the next two images.

SOA $a_{cdm}(443)$ (m^{-1})



SOA $a_{cdm}(443)$ (m^{-1})

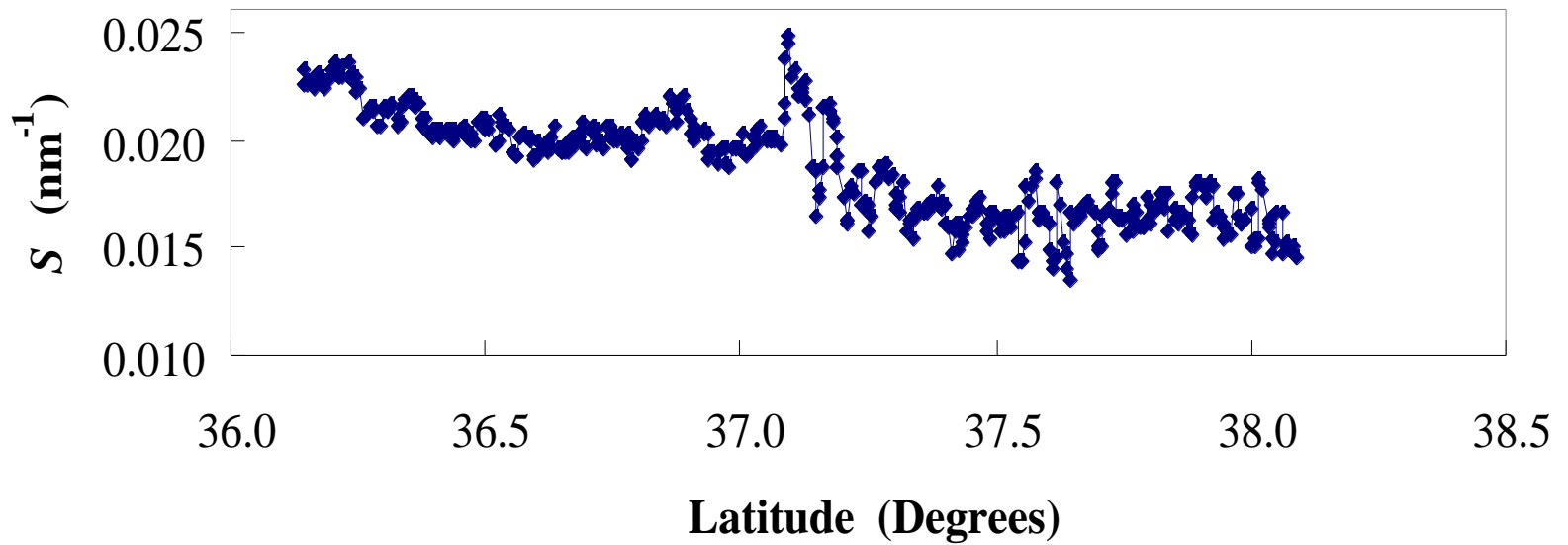




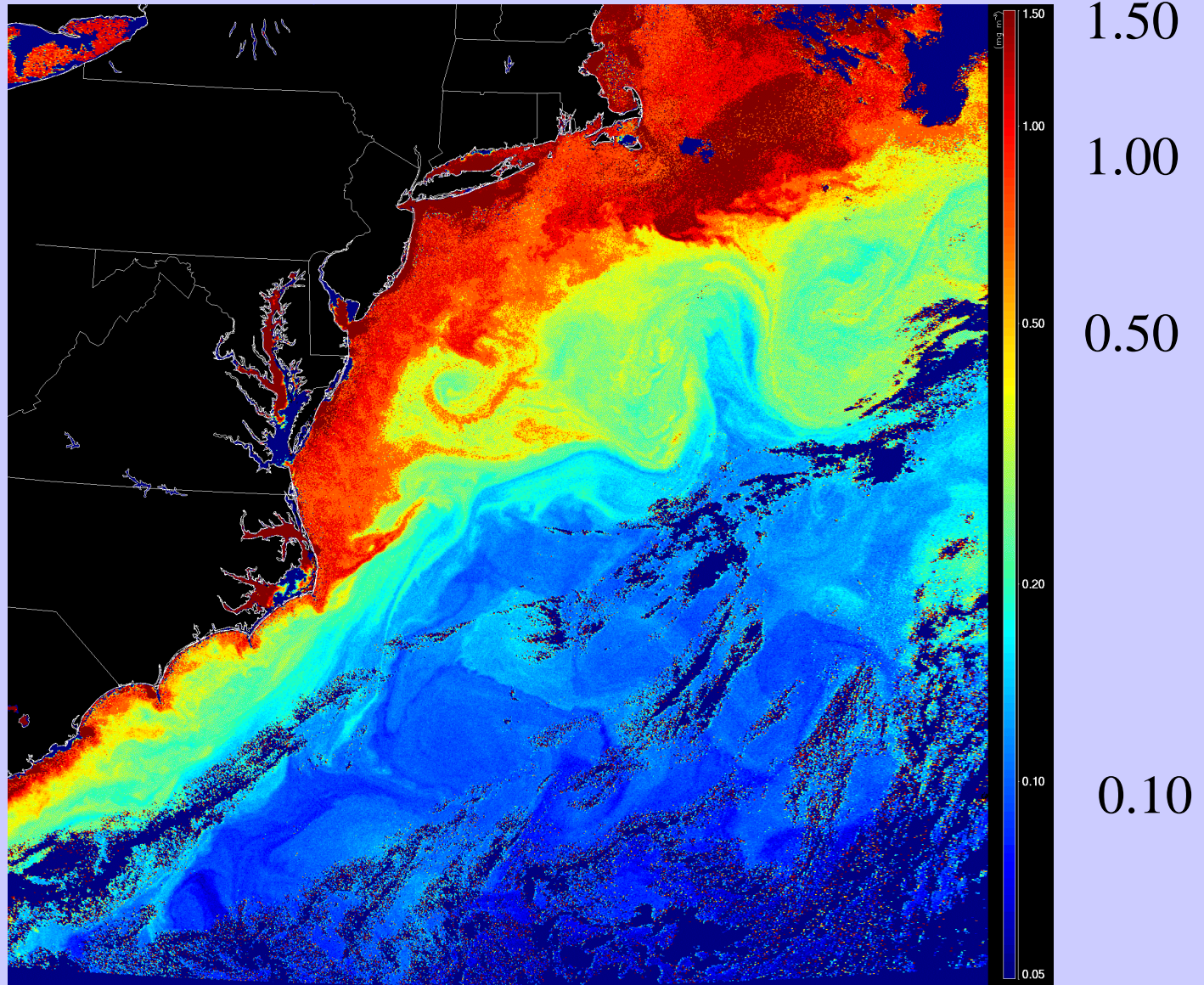
Comparison of SOA and AOL $a_{cdm}(443)$
along the North-South Track

The value of S required to bring the SOA retrieved $a_{cdm}(443)$ into confluence with the AOL-retrieved $a_{CDOM}(443)$ at each point along the track the track was determined and shown in the next slide. The resulting S values show a clear trend of decreasing into the mesotrophic waters as would be expected (Green and Blough, 1994). Similar results are found for the other two tracks.

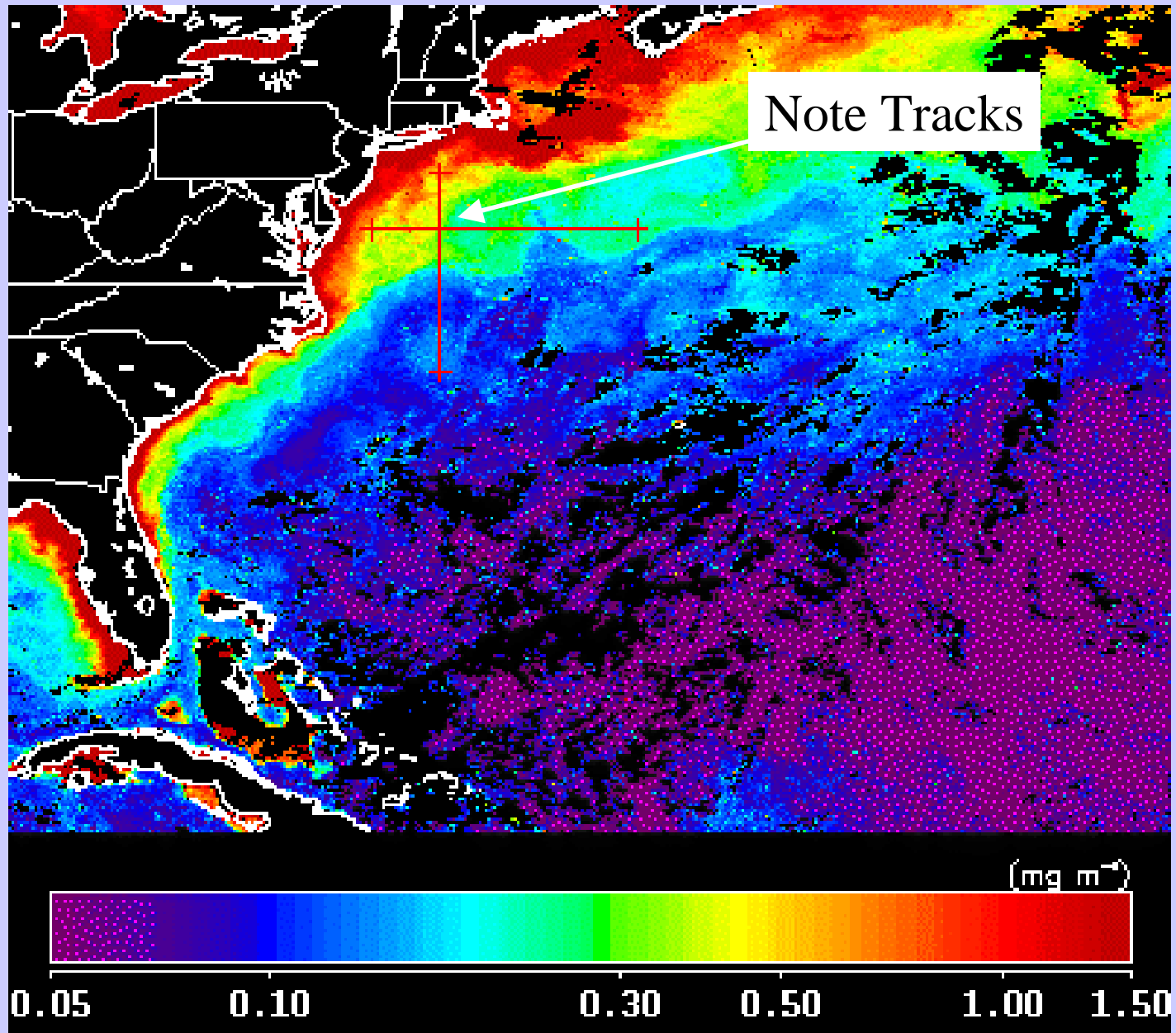
**Required "S" for Exact
AOL-SOA Agreement
Along North-South Track**



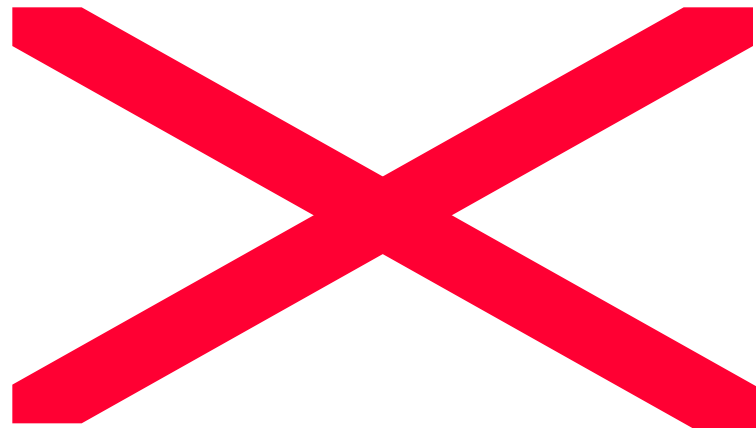
SOA *Chl a* (mg/m³)



SeaWiFS 8-day mean *Chl a*



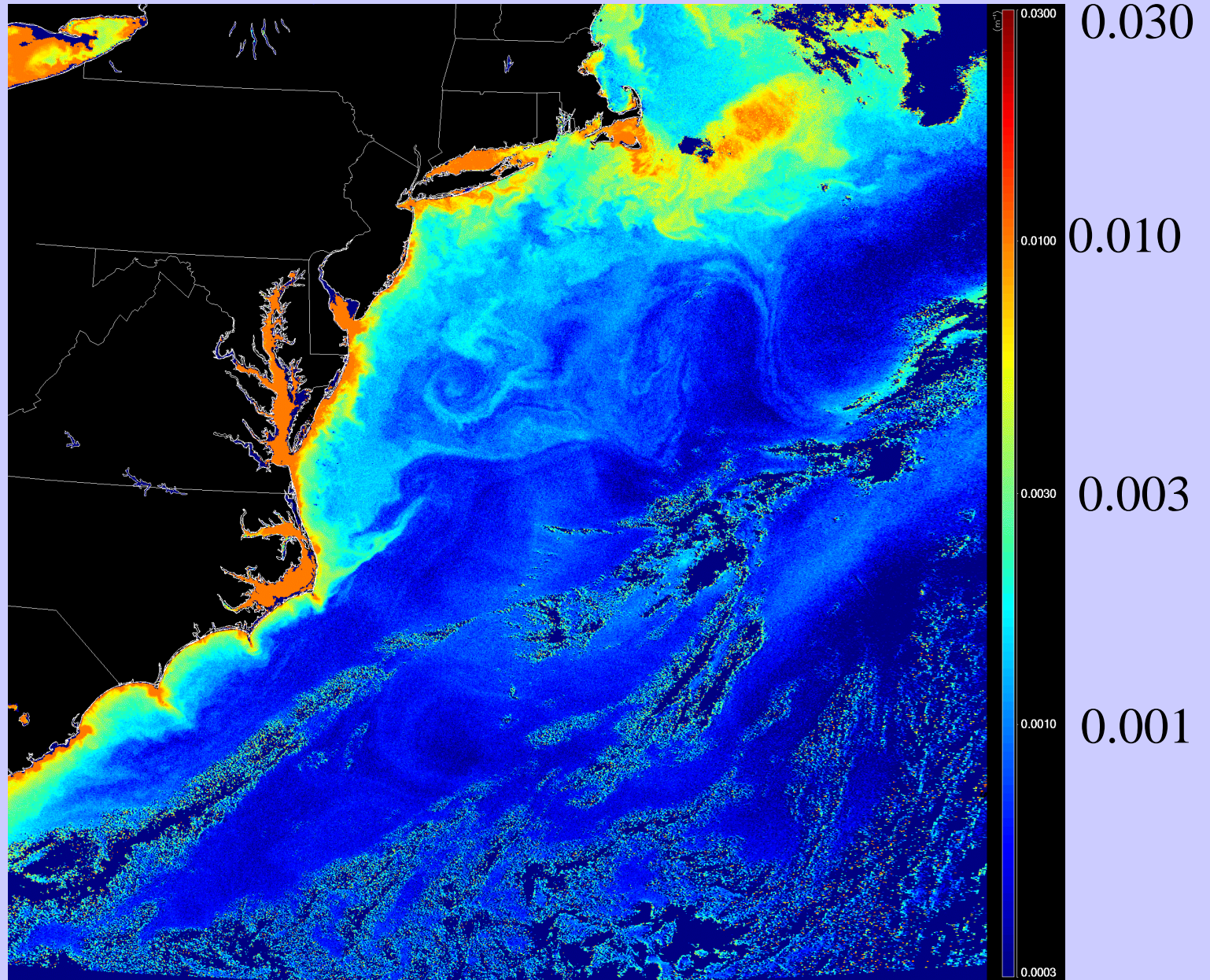
Comparison with SeaWiFS



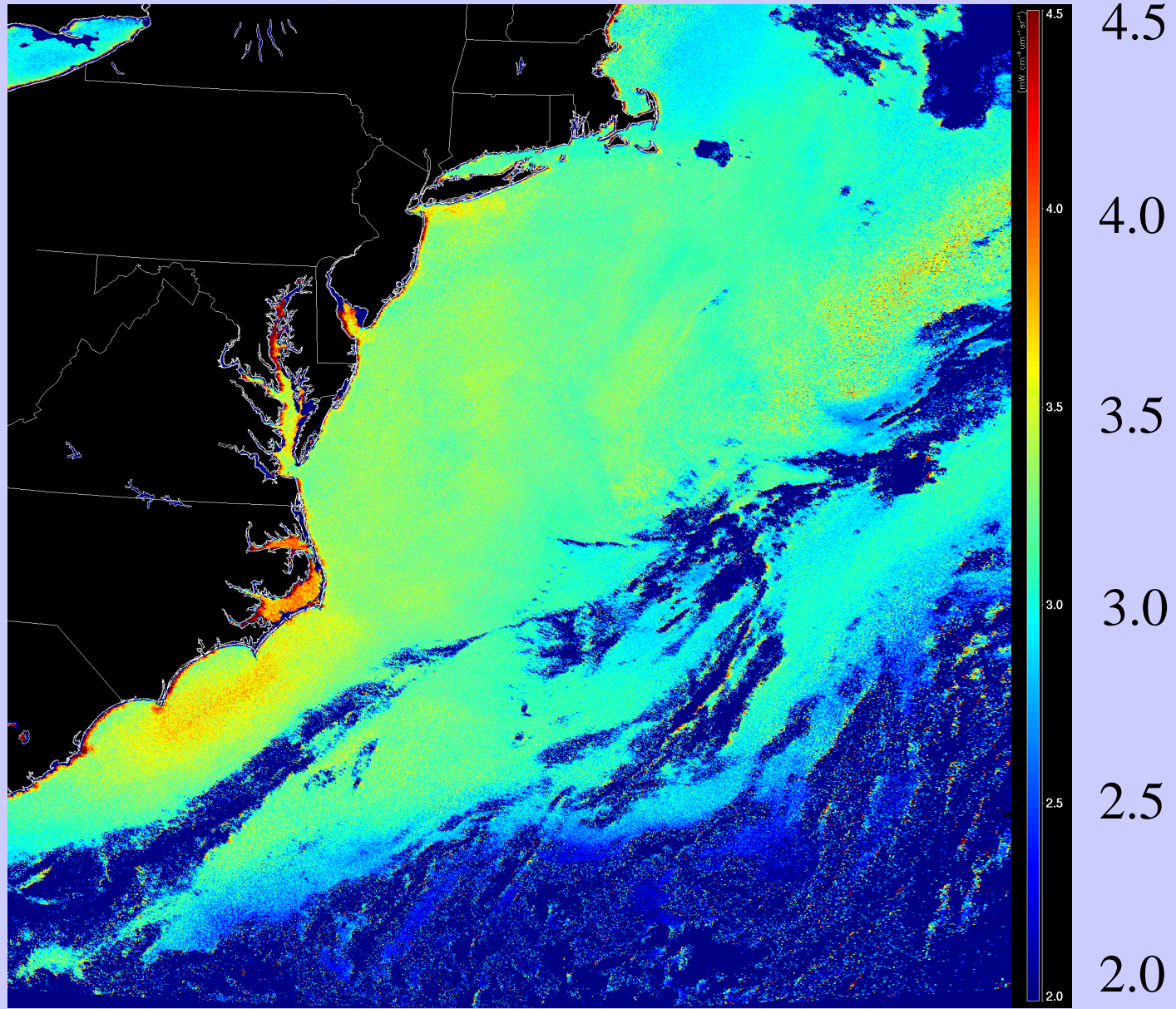
Comparison with SeaWiFS



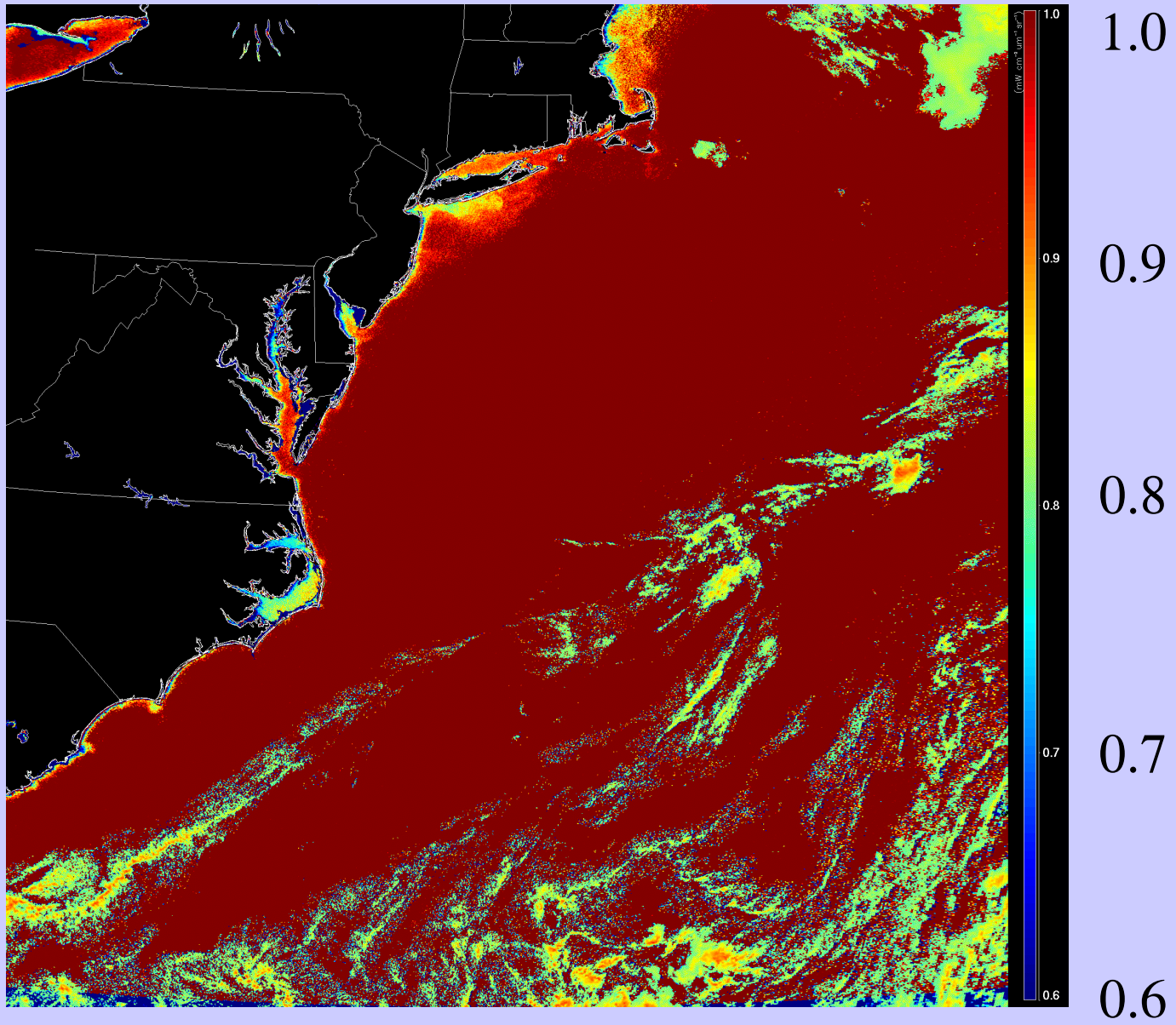
$b_{bp}(443) \text{ (m}^{-1}\text{)}$



V



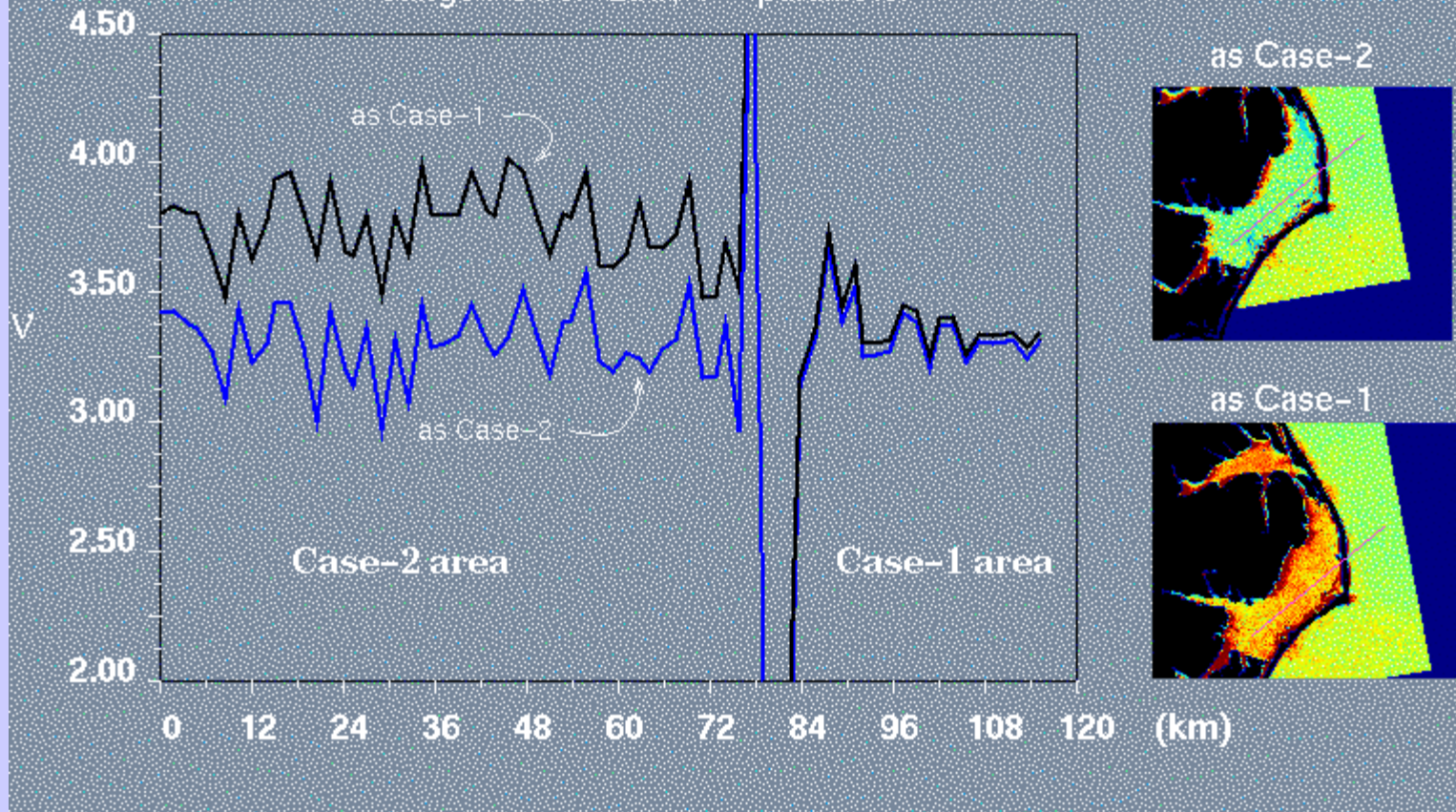
ω_0



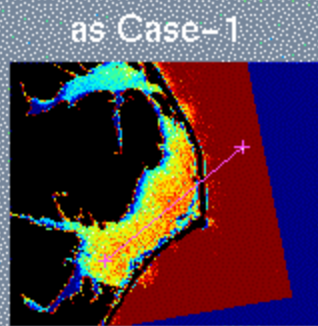
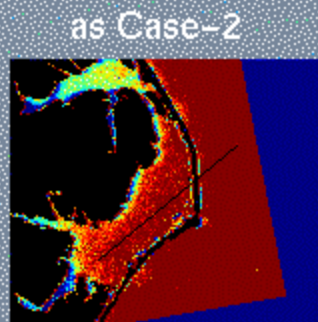
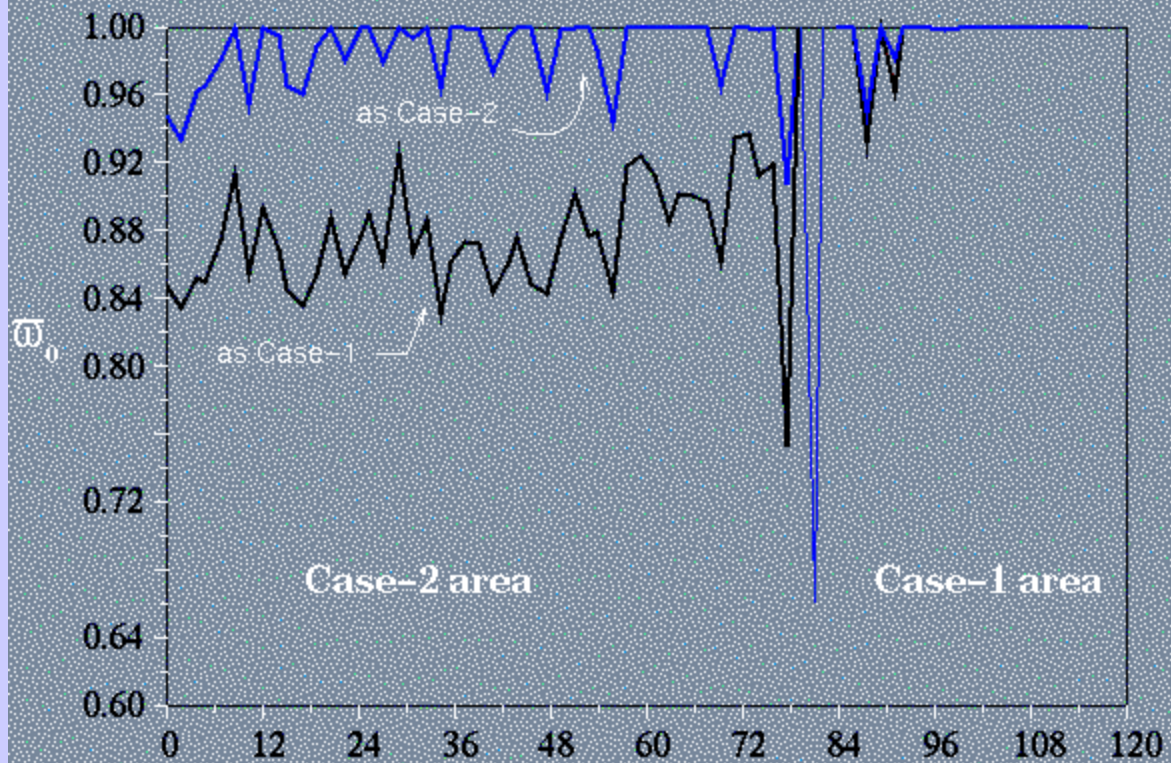
Extension to Case 2 Waters

- In Case 2 waters, we operate the algorithm as in Case 1 (waters), i.e., assuming that $\rho_w(\lambda_{NIR}) = \rho_v(\lambda_{NIR}) + \rho_a(\lambda_{NIR})$
 - Then we use the retrieved values of C , $a_{cdm}(443)$, and $b_{bp}(443)$ to provide an estimate of $\rho_w(\lambda_{NIR})$ in the NIR, and the retrieved values of v , τ_a , m_r , and m_i to estimate t_v and t_s and the NIR.
 - These estimates are subtracted from the total, i.e.,
 - The $v - \tau_a$ portion of the algorithm is then operated with
- instead of $\rho_t(\lambda_{NIR})$, to estimate new constraints $v(m_r, m_i)$ and $\tau_a(m_r, m_i)$, and to initiate a new optimization, etc.

Comparison between Case2 and non-Case2 processing
Junge Power-Law, V - parameter



Comparison between Case2 and non-Case2 processing
single-scattering albedo, w_0

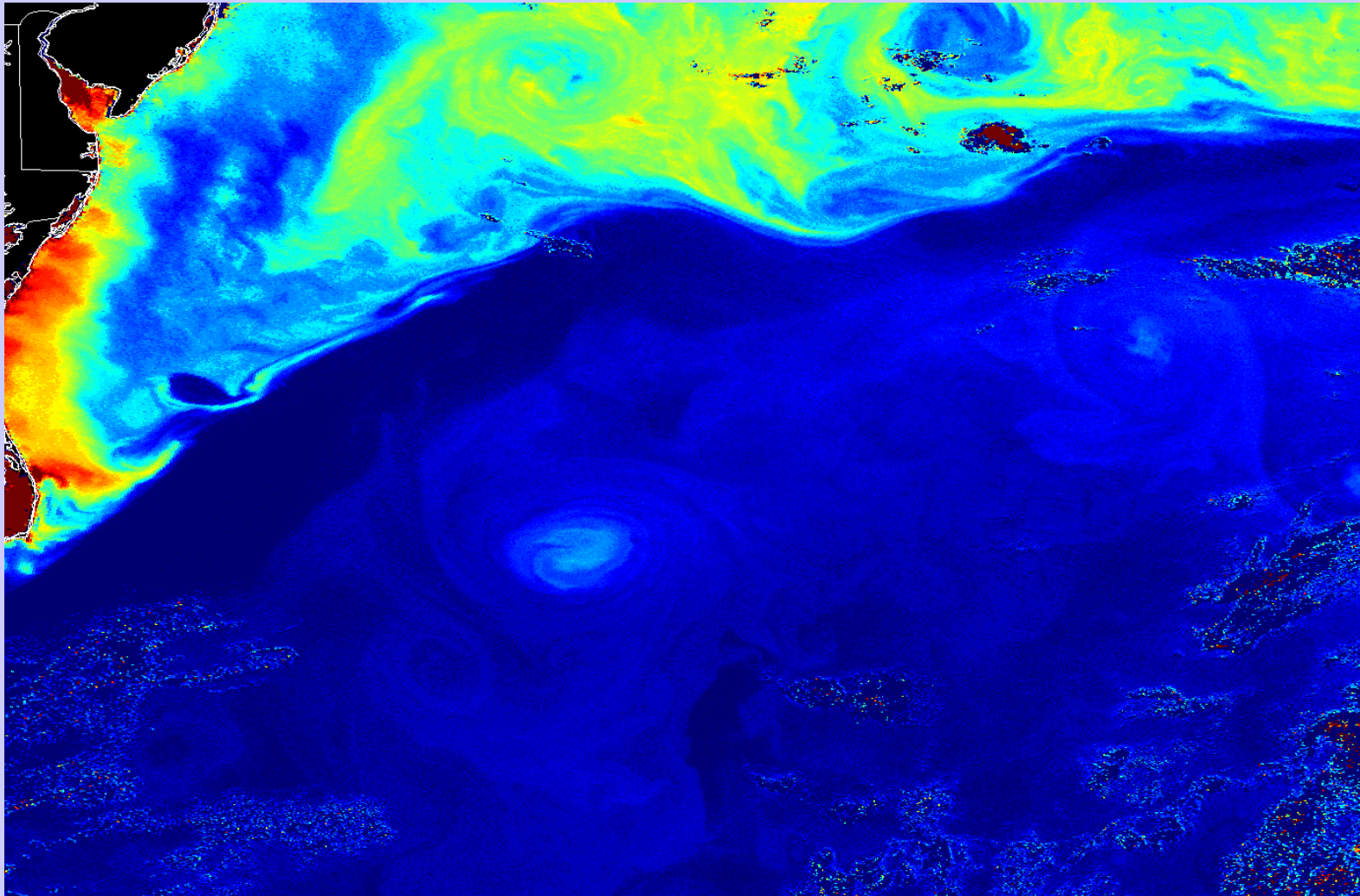


Incorporation into the MODIS Code : A Status Report

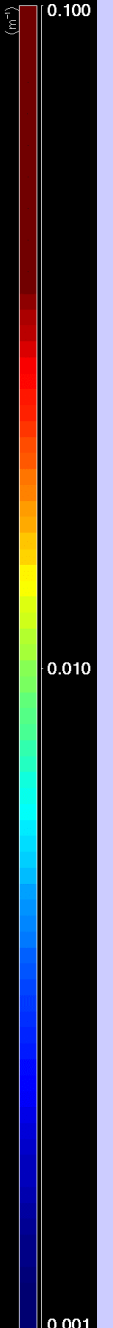
Processing philosophy

- Spectral Optimization Algorithm is slow, so at present we must restrict application to sub-granulars.
- Unlike the Case 1 $\rho_w(\lambda)$ model, the Case 2 $\rho_w(\lambda)$ model will most likely be site specific, i.e., the parameters in the GS97 model $\{a_{ph0}(\lambda), S, \text{ and } n\}$ will depend on the target location.
- Our goal is to provide processing code that can be used for any location, given model parameters for that location. Individual investigators must supply $a_{ph0}(\lambda)$, S , and n .

$$b_{bp} \text{ (m}^{-1}\text{)}$$

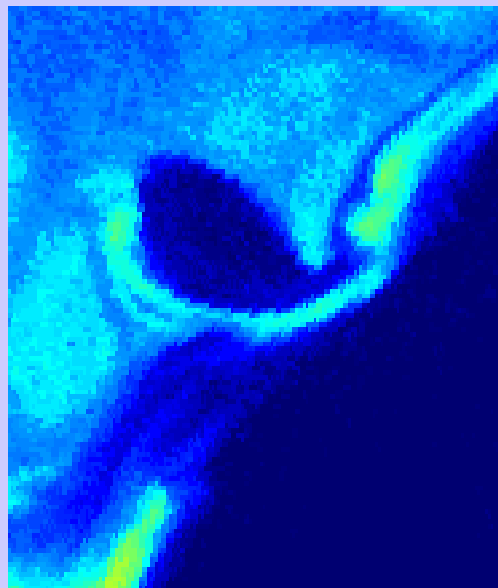


SeaWiFS

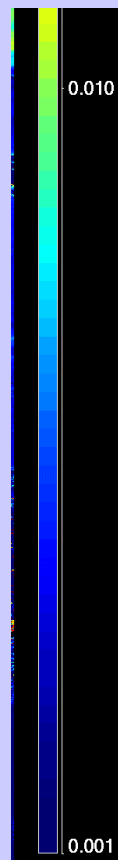
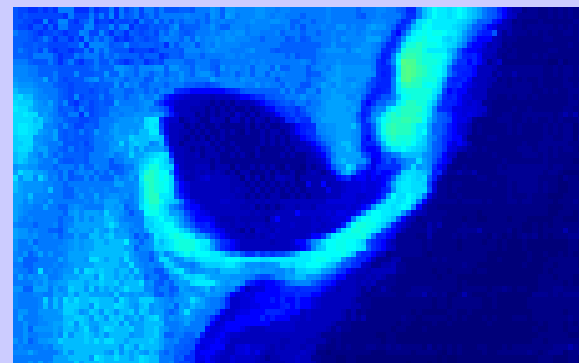


$$b_{bp} \text{ (m}^{-1}\text{)}$$

SeaWiFS

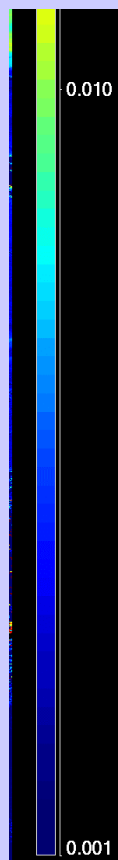
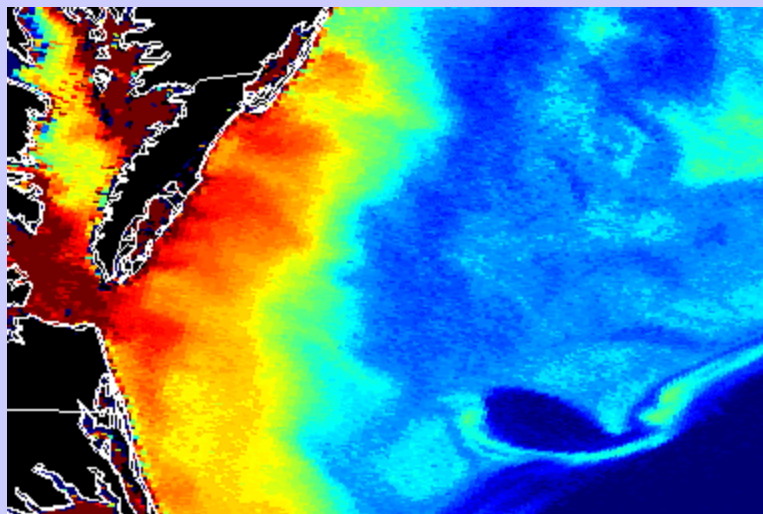


MODIS

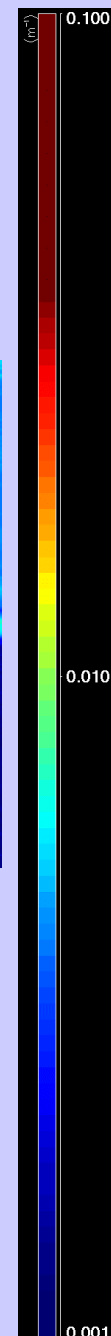
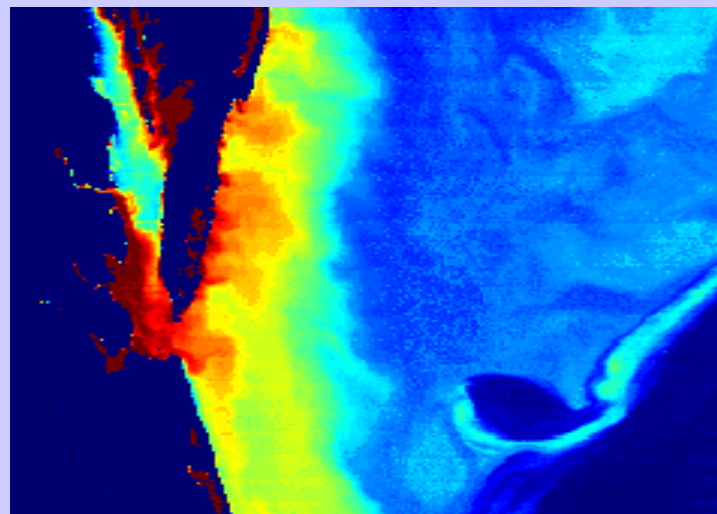


$$b_{bp} \text{ (m}^{-1}\text{)}$$

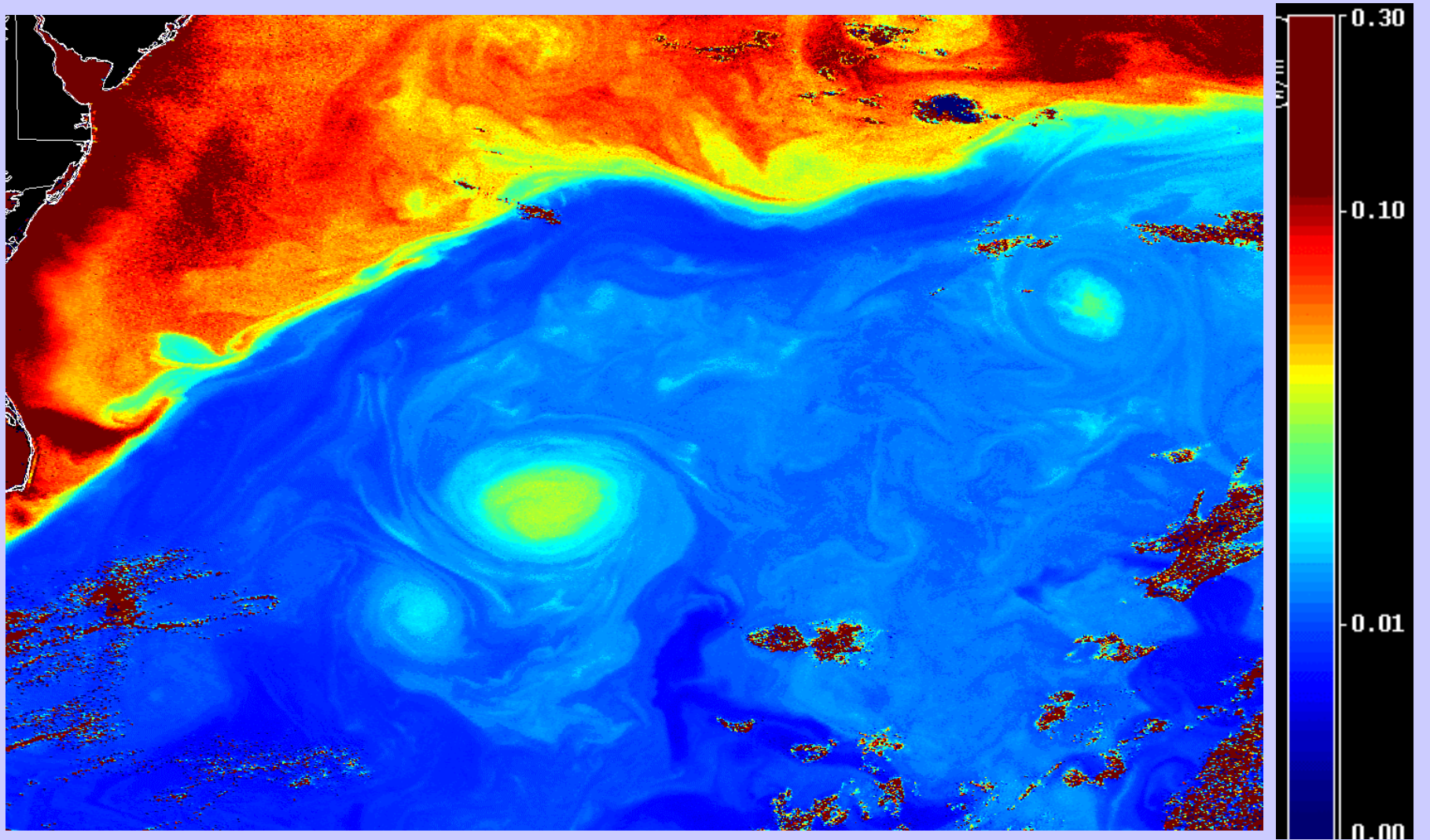
SeaWiFS



MODIS



$$a_{cdm} \text{ (m}^{-1}\text{)}$$

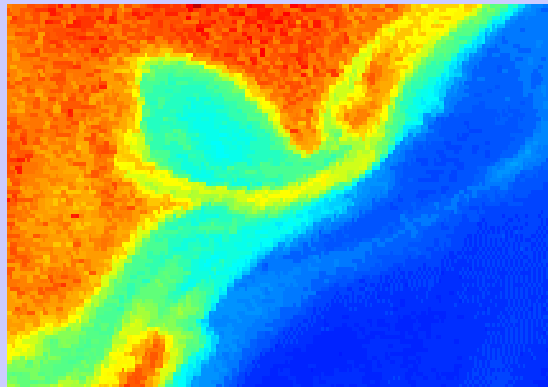


SeaWiFS

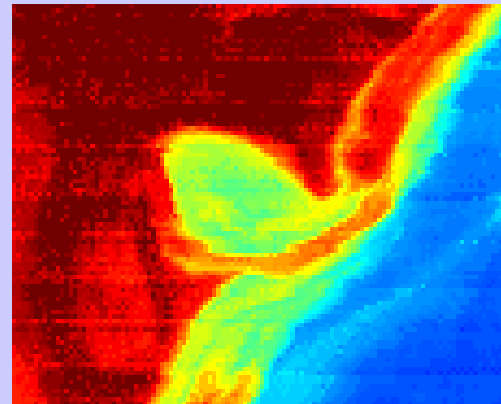
0.003

$$a_{cdm} \text{ (m}^{-1}\text{)}$$

SeaWiFS



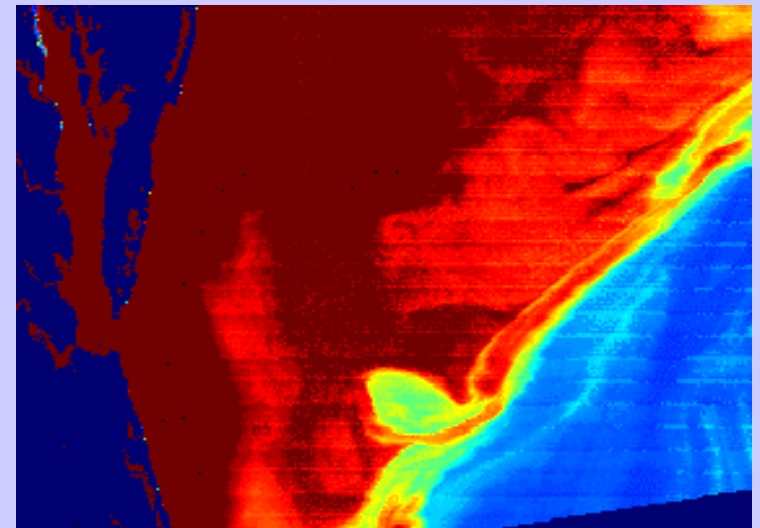
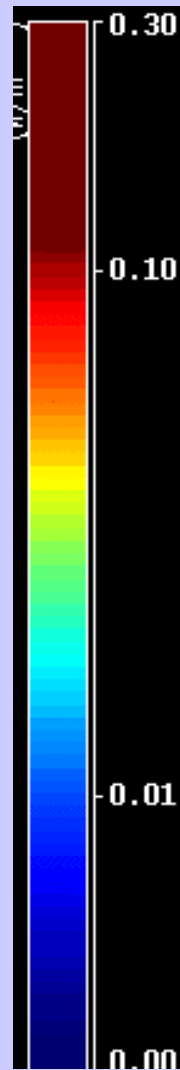
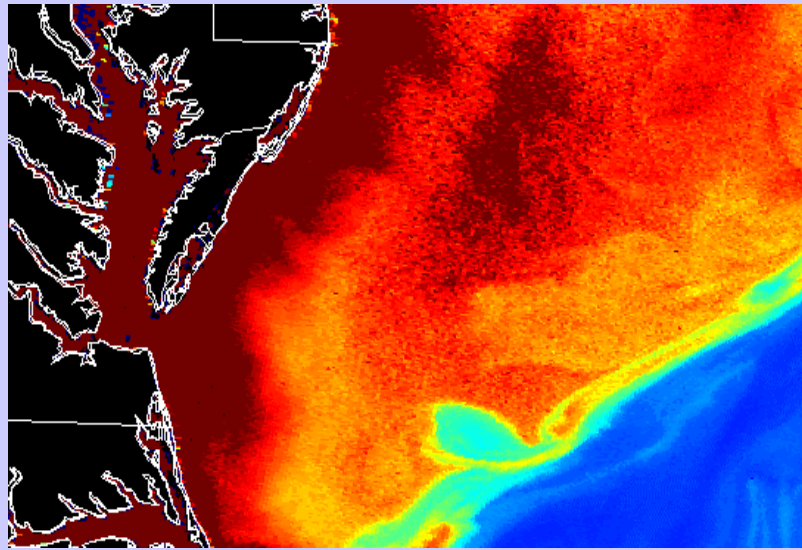
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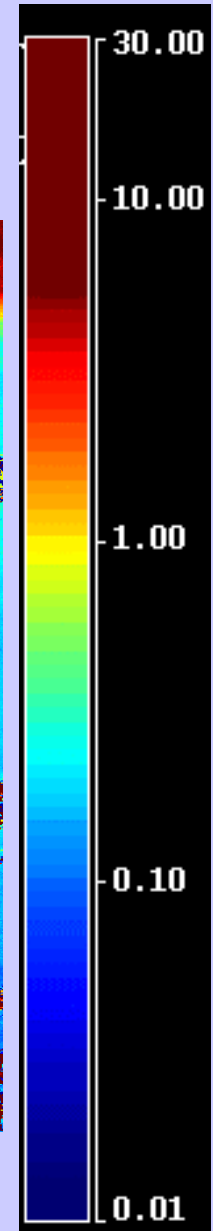
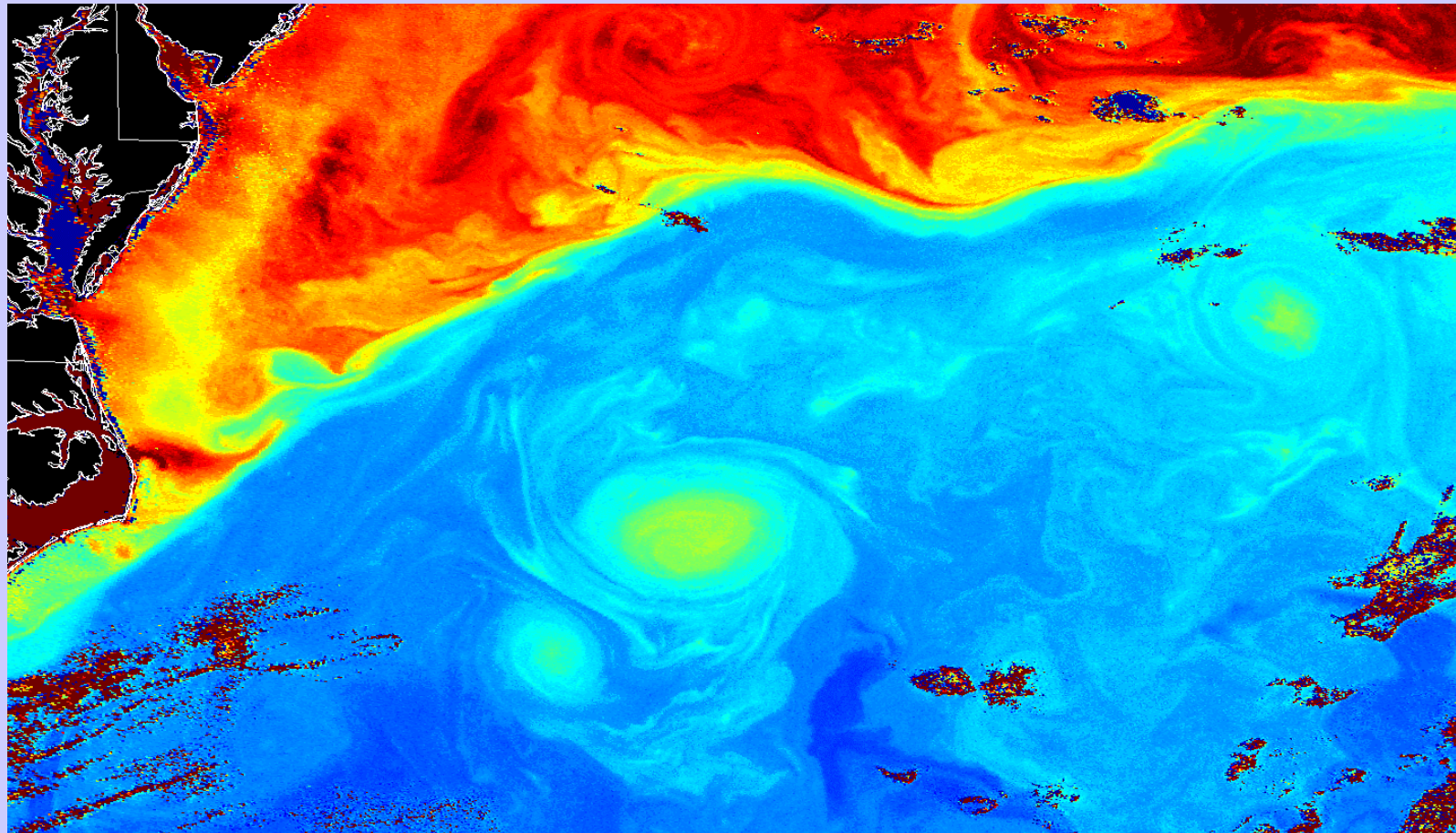
$$a_{cdm} \text{ (m}^{-1}\text{)}$$

SeaWiFS

MODIS



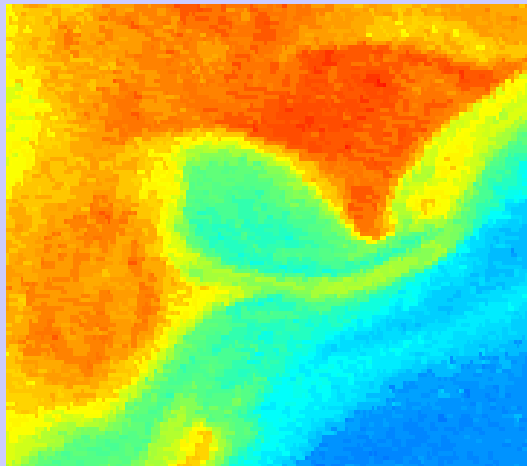
Chl (mg m⁻³)



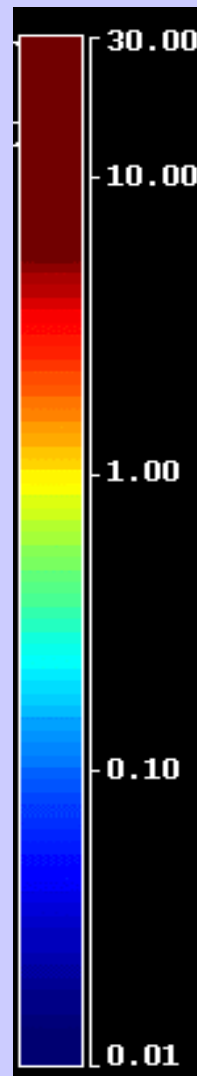
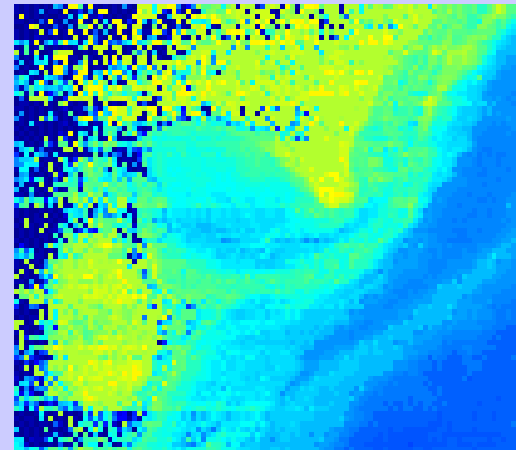
SeaWiFS

Chl (mg m⁻³)

SeaWiFS

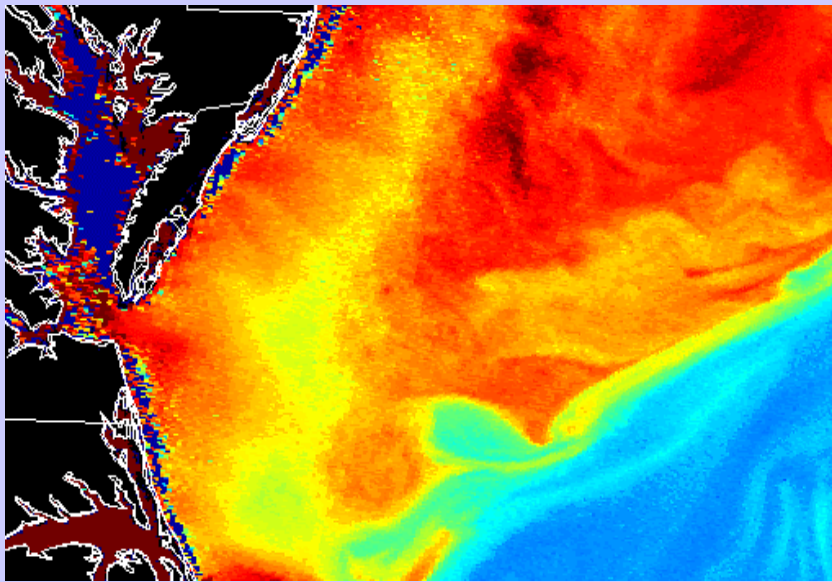


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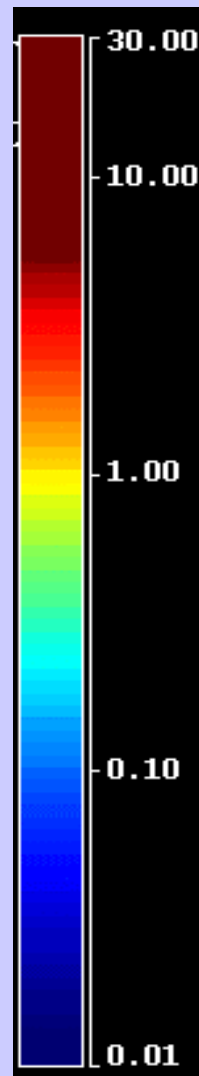
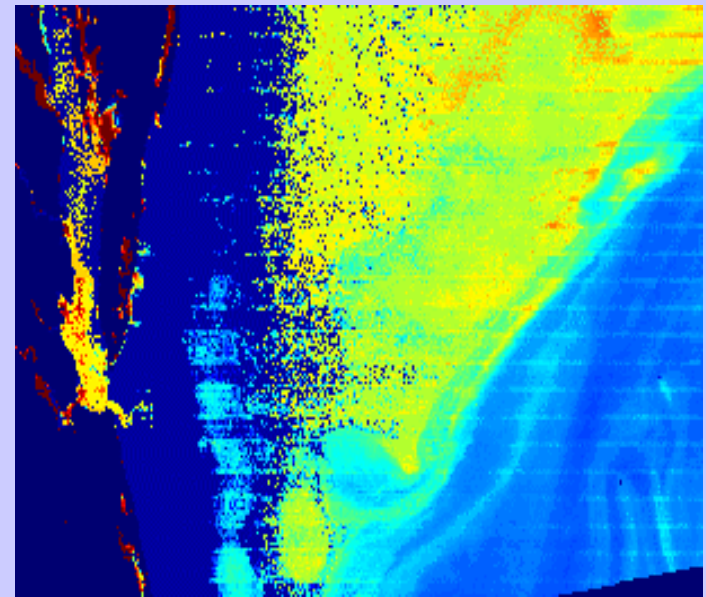


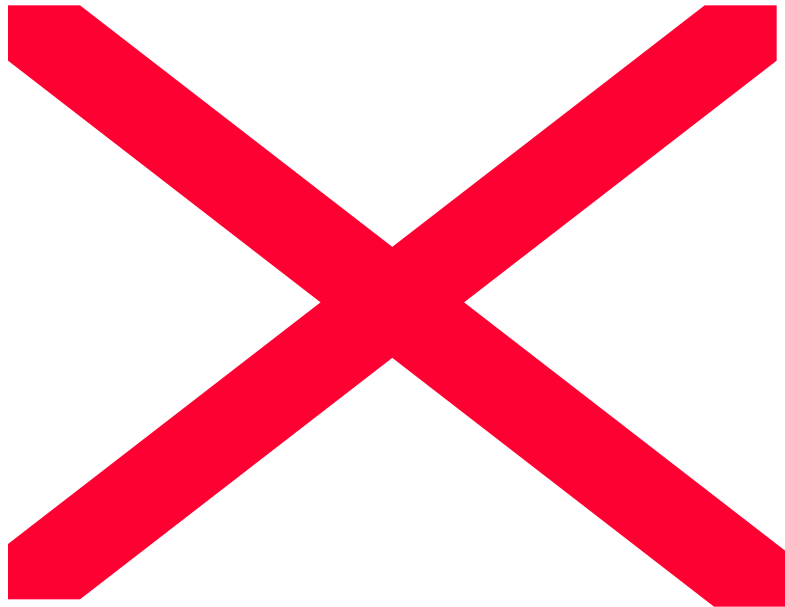
Chl (mg m⁻³)

SeaWiFS



MODIS





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